

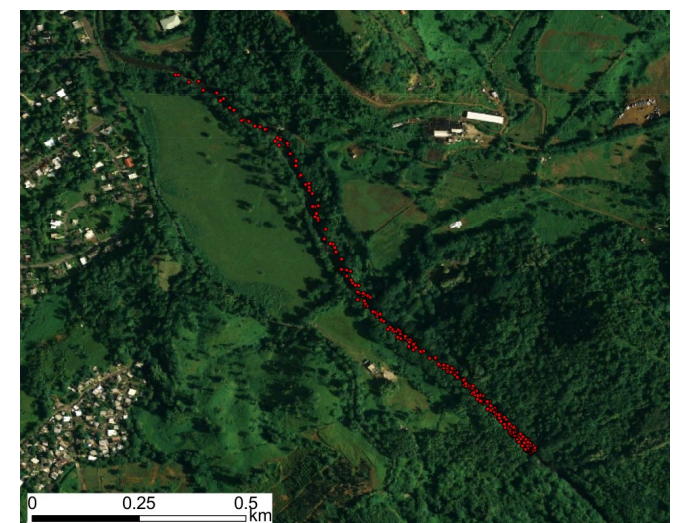
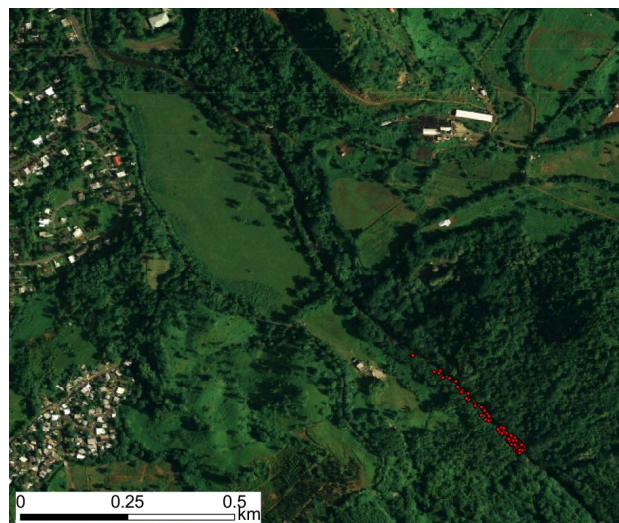
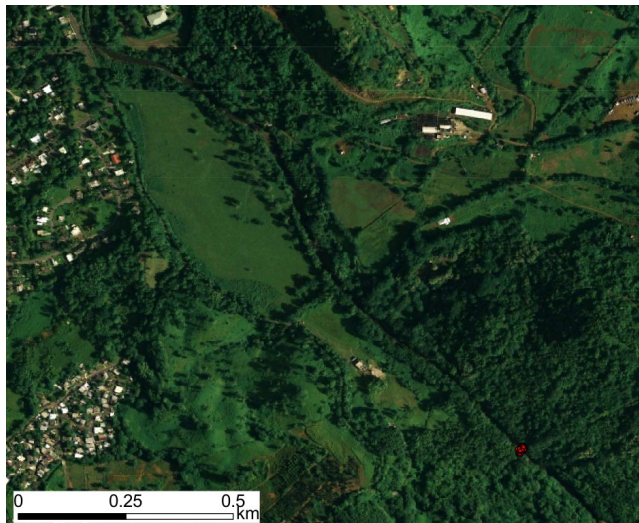
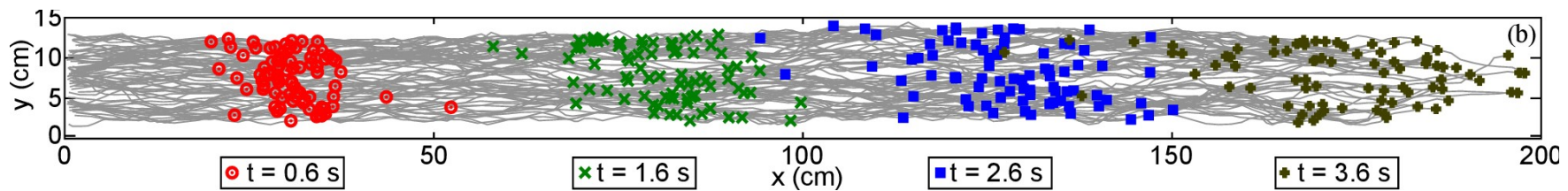
# Physical causes and modeling challenges of anomalous diffusion of sediment tracers

**Douglas Jerolmack**

Earth & Environmental  
Science, UPenn

[[sediment@sas.upenn.edu](mailto:sediment@sas.upenn.edu)]

“Bridging the Gap”,  
Princeton U., 2 Oct. 2012



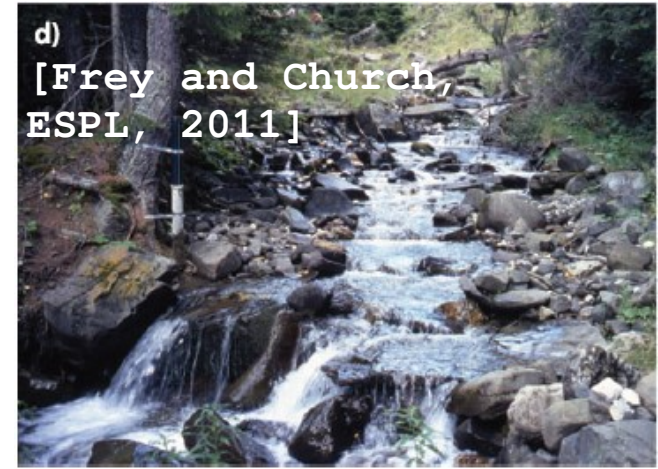
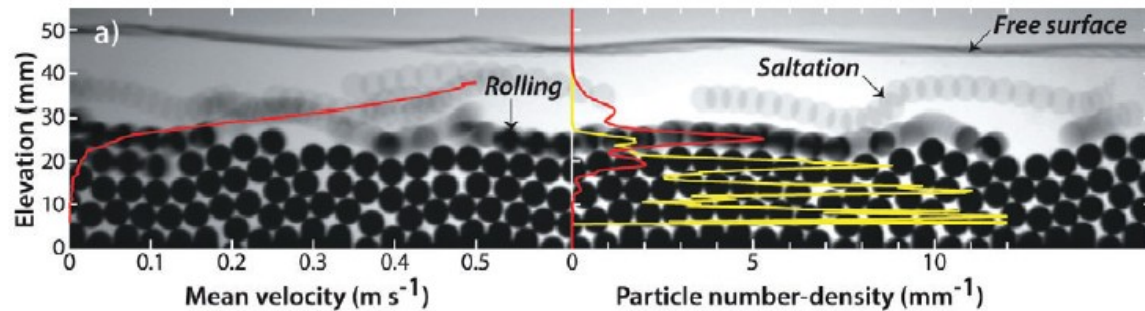
# “Bed load” transport: multi-phase flow problem

Grains supported by and in frequent contact with the bed.

A well defined “layer” describable by:

1. Particle volume,  $\delta v$  [ $L^3$ ].
2. Average velocity,  $u_s$ , of bed load sediment [ $L/T$ ].
3. Surface density,  $n$ , of moving particles [ $\#/L^2$ ].

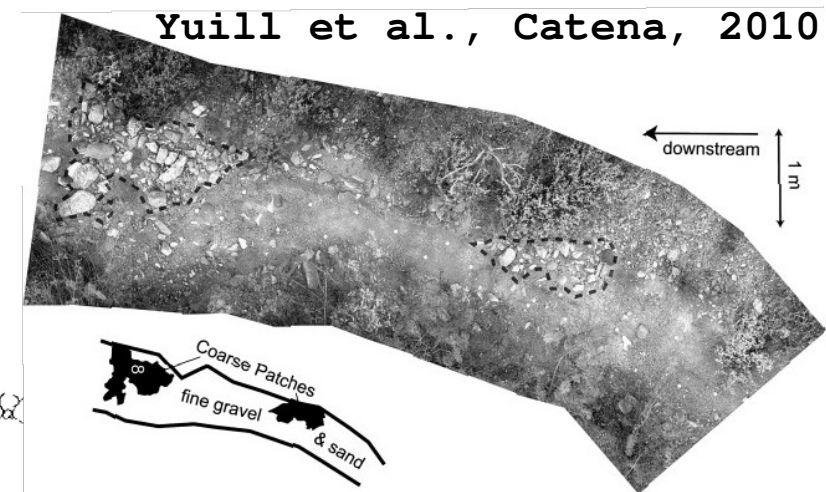
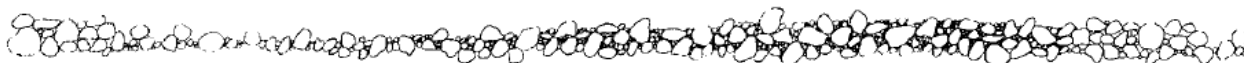
$$q_s = \delta v n u_s$$



Grains clump together - **bed forms**

- “Force chains” form and break
- Slow, stochastic shearing of granular bed
- Spatial sorting and clusters

Side view

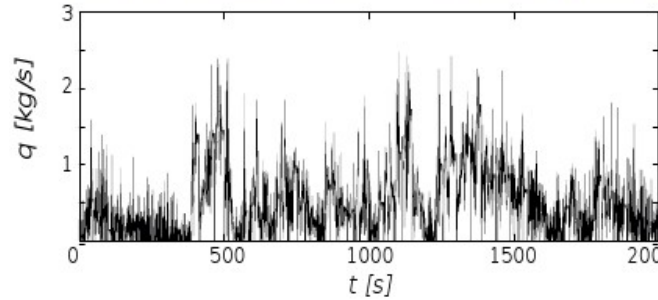


# Near threshold transport: "avalanching" and spatially heterogeneous dynamics (?)



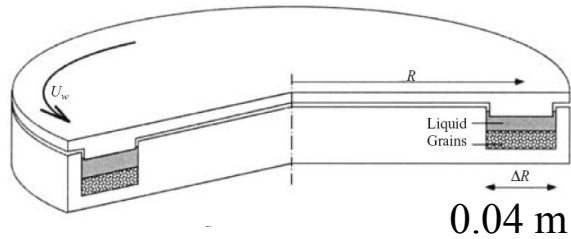
← 3 m →

[Singh et al., JGR-ES, 2009]



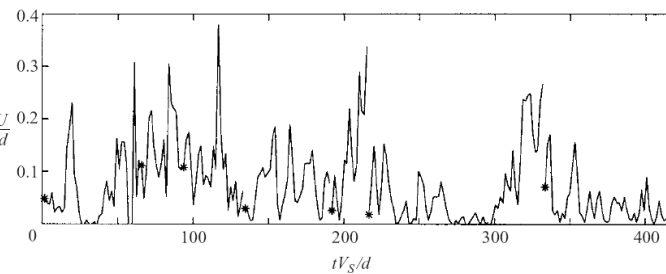
Turbulence, granular collisions, grain size dispersion.

Intermittent transport under steady flow.



0.04 m

[Charru et al., JFM, 2004]



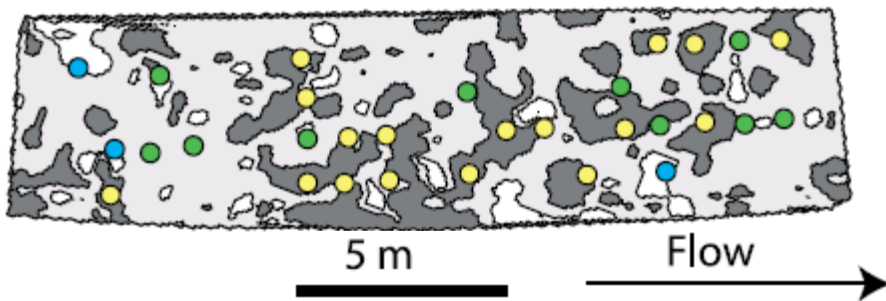
Laminar flow, uniform beads.

Intermittent transport → collective grain motion.

JOURNAL OF GEOPHYSICAL RESEARCH, VOL. 117, F01014, doi:10.1029/2011JF002120, 2012

Using multiple bed load measurements: Toward the identification of bed dilation and contraction in gravel-bed rivers

G. A. Marquis<sup>1</sup> and A. G. Roy<sup>1</sup>

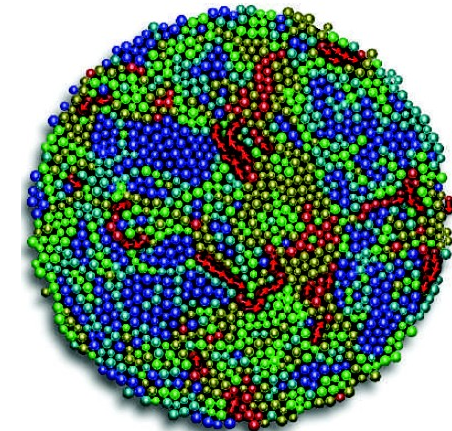


- Elevation increase
- Stable elevation
- Elevation decrease
- 2A. Dilation
- 3A. Contraction
- 1B. Active bed / no elevation change

nature physics | VOL 3 | APRIL 2007 | www.nature.com/naturephysics

Measurement of growing dynamical length scales and prediction of the jamming transition in a granular material

AARON S. KEYS<sup>1\*</sup>, ADAM R. ABATE<sup>2\*</sup>, SHARON C. GLOTZER<sup>1,3†</sup> AND DOUGLAS J. DURIAN<sup>2†</sup>



Increasing mobility →

# One solution: CFD coupled to DEM

*Sedimentology* (2003) 50, 279–301

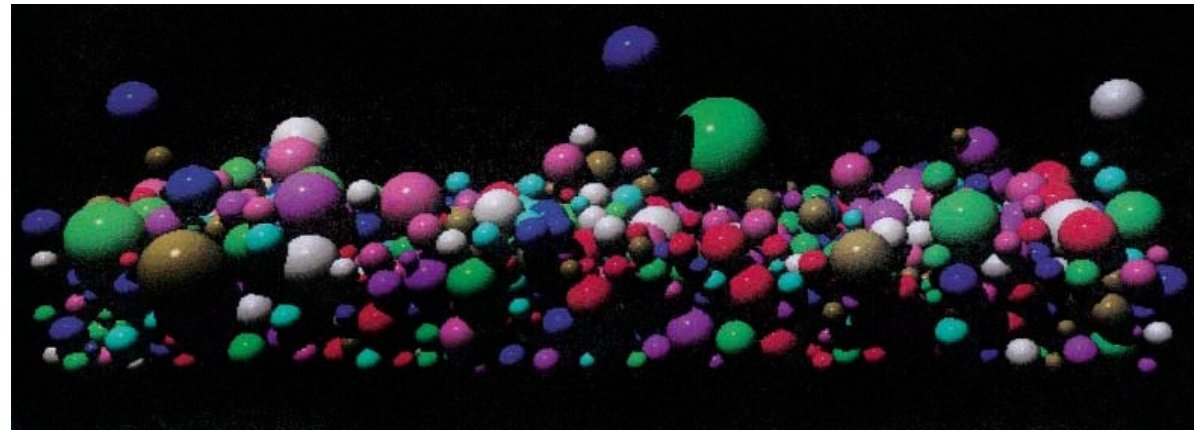
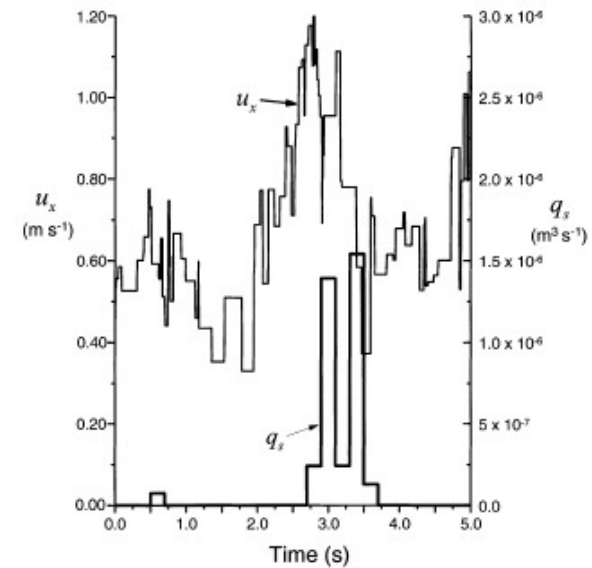
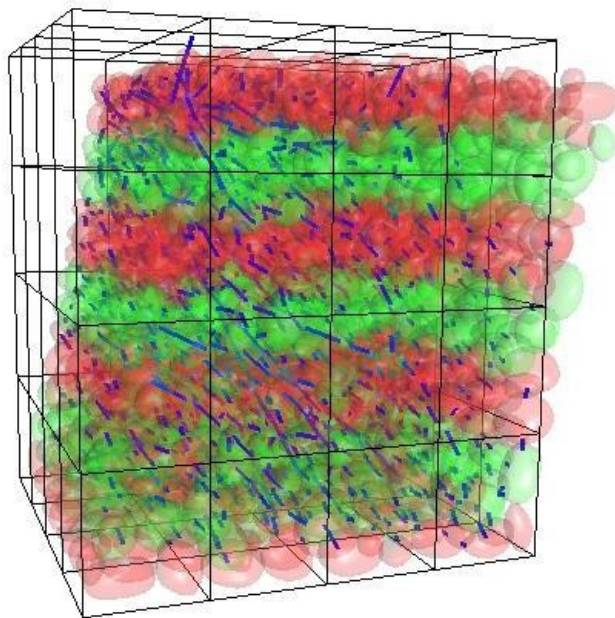
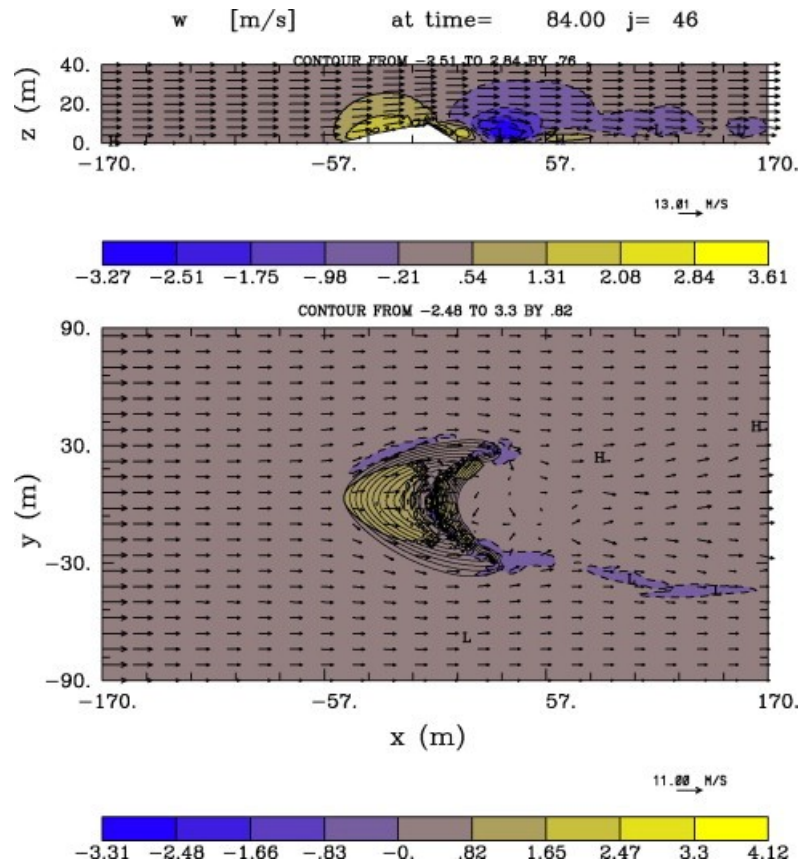
## Direct numerical simulation of bedload transport using a local, dynamic boundary condition

MARK W. SCHMEECKLE\* and JONATHAN M. NELSON†

## Numerical simulation of turbulent sediment transport, from bed load to saltation.

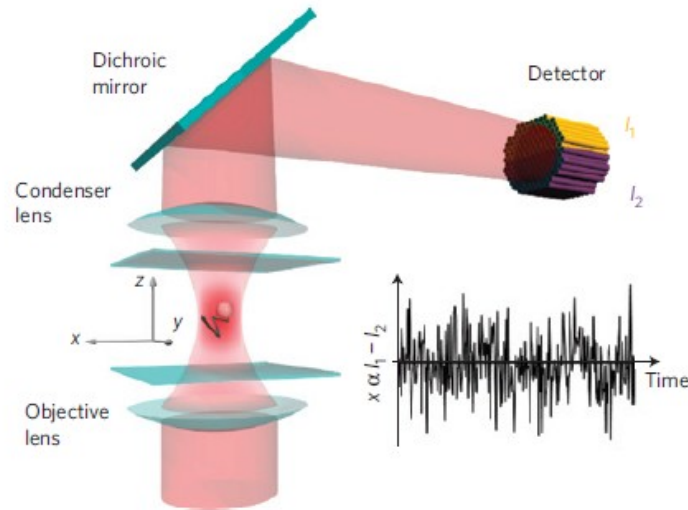
Orencio Durán,<sup>1, a)</sup> Bruno Andreotti,<sup>1</sup> and Philippe Claudin<sup>1</sup>

arXiv:1111.6898v2 [cond-mat.soft] 29 May 2012



## Direct observation of the full transition from ballistic to diffusive Brownian motion in a liquid

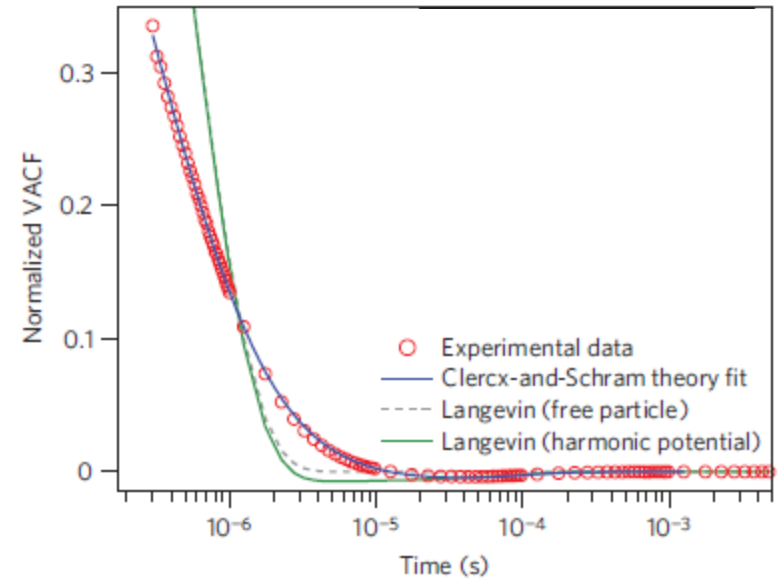
Rongxin Huang<sup>1</sup>, Isaac Chavez<sup>1</sup>, Katja M. Taute<sup>1</sup>, Branimir Lukić<sup>2</sup>, Sylvia Jeney<sup>2</sup>, Mark G. Raizen<sup>1</sup> and Ernst-Ludwig Florin<sup>1\*</sup>



**Figure 1 | Schematic diagram of the experiment.** A single micrometre-size particle in water is undergoing Brownian motion in the observation volume given by an optical trap.

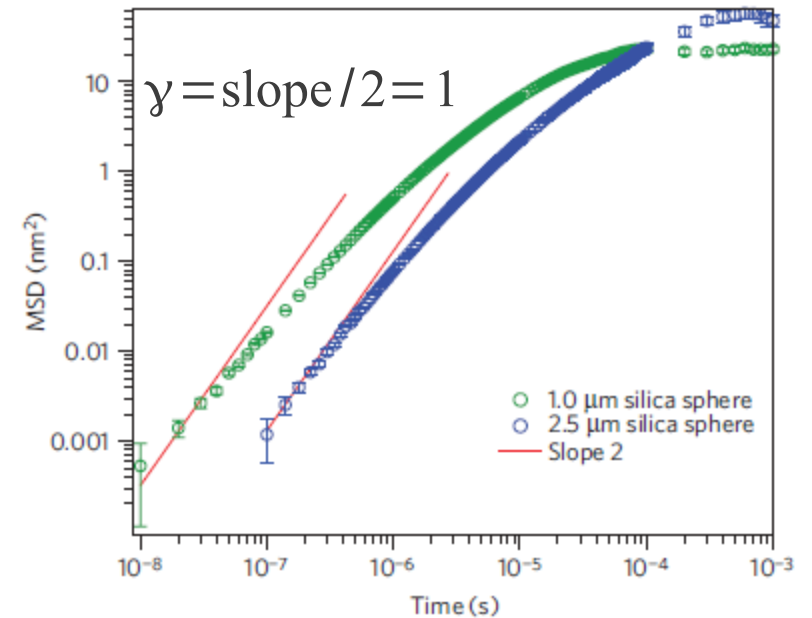
$$T_p = m/b \sim 10^{-6} \text{ s}$$

Inertial timescale      Particle mass      Stokes drag



**Figure 4 | Experimental VACF and theoretical description.** The VACF

Ballistic transport at short time

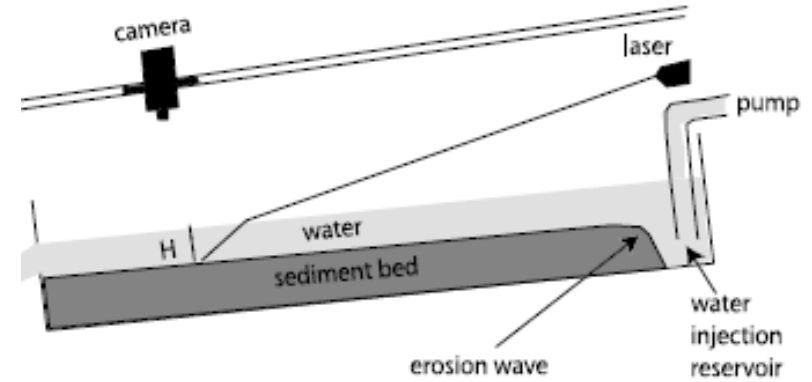


**Figure 2 | Example MSD for silica particles 1 µm and 2.5 µm in diameter.**

# Bed load: Brownian motion with drift?



**PennSeD**  
UNIVERSITY OF PENNSYLVANIA SEDIMENT DYNAMICS LAB



(a)

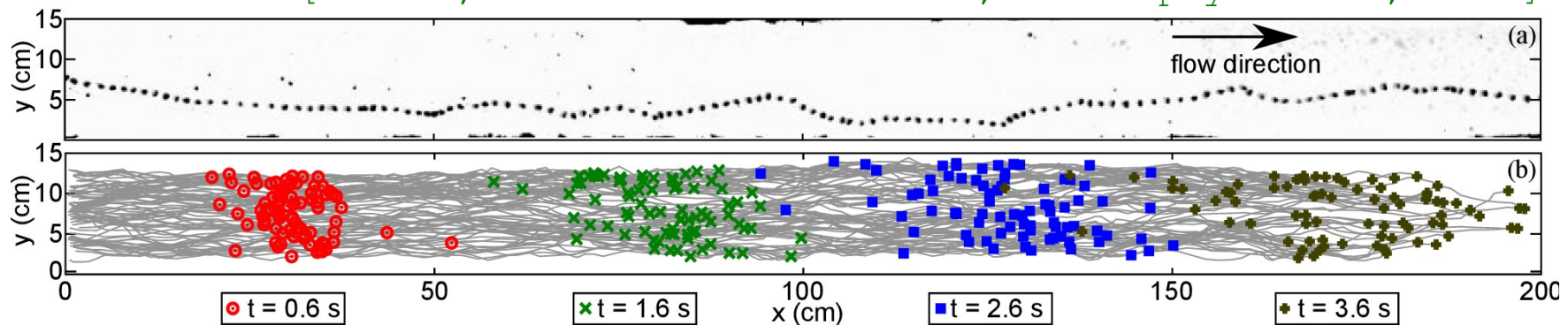
Momentum balance for particle in a turbulent shear flow:

$$T_p = \frac{4\rho_s D}{3\rho C_d} \left( \frac{1}{u_f - v_x} \right) \sim 10^{-1} \text{ s}$$

Drag,  
inviscid  
limit

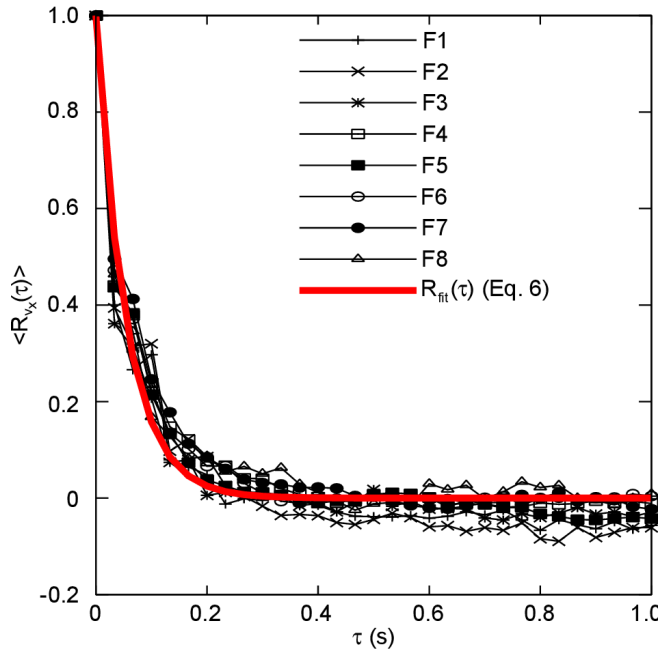
Slip velocity

[Martin, Jerolmack and Schumer, J. Geophys. Res., 2012]

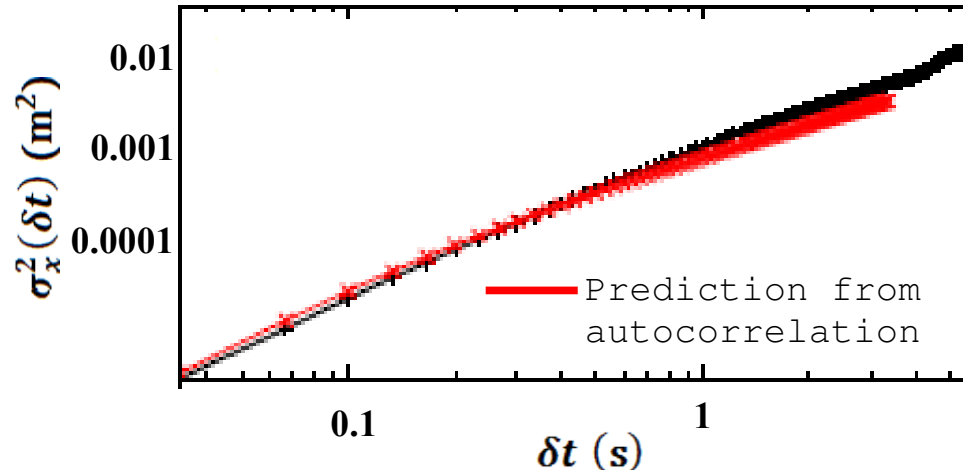


## Velocity autocorrelation:

Controlled by inertia



## Dispersion: inertial at short time



**Short timescales: yes, diffusive particle transport:**

**Statistical mechanics** may be used to derive macroscopic transport laws from stochastic particle motions.

A probabilistic description of the bed load sediment flux:

### 1. Theory

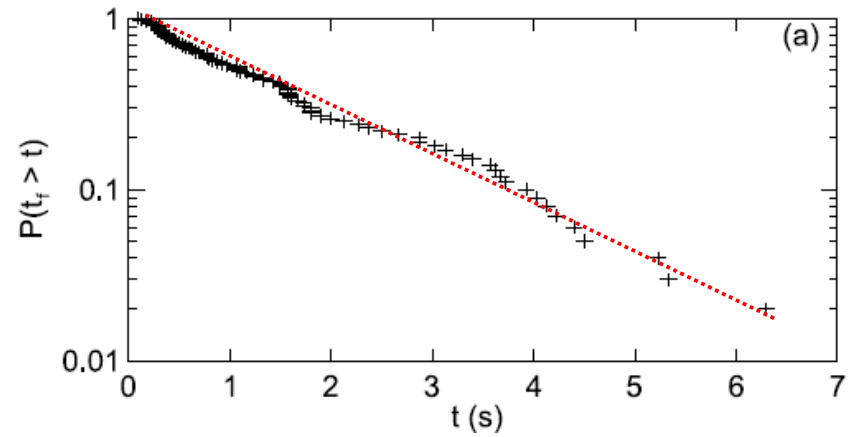
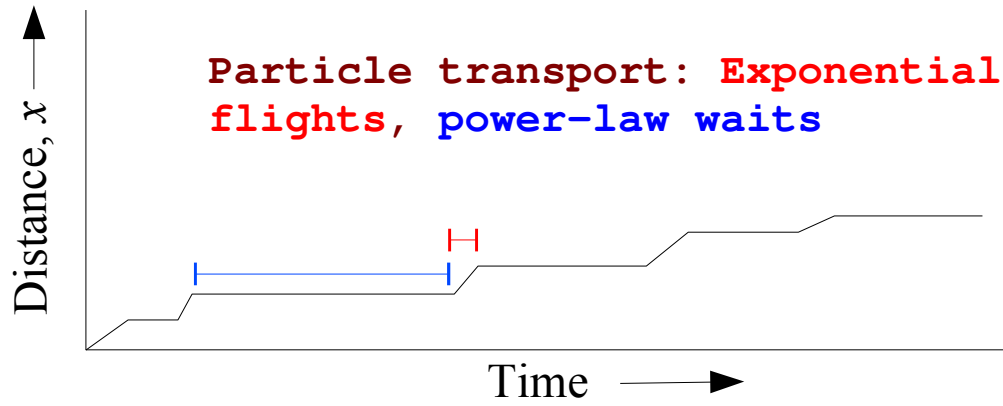
David Jon Furbish,<sup>1</sup> Peter K. Haff,<sup>2</sup> John C. Roseberry,<sup>1</sup> and Mark W. Schmeckle<sup>3</sup>

A probabilistic description of the bed load sediment flux:

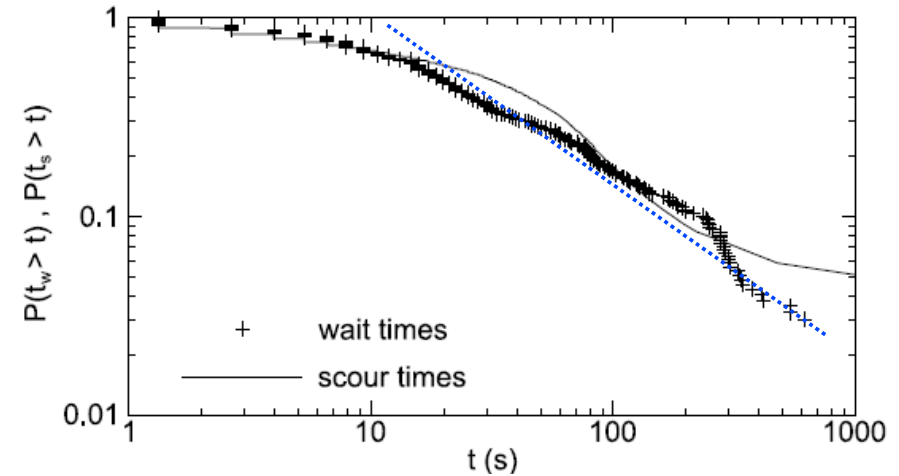
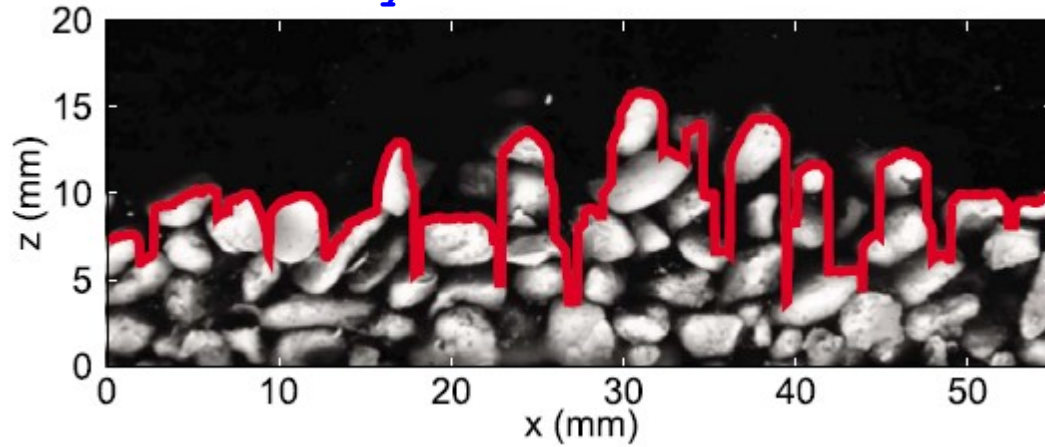
### 4. Fickian diffusion at low transport rates

David Jon Furbish,<sup>1</sup> Ashley E. Ball,<sup>2</sup> and Mark W. Schmeckle<sup>3</sup>

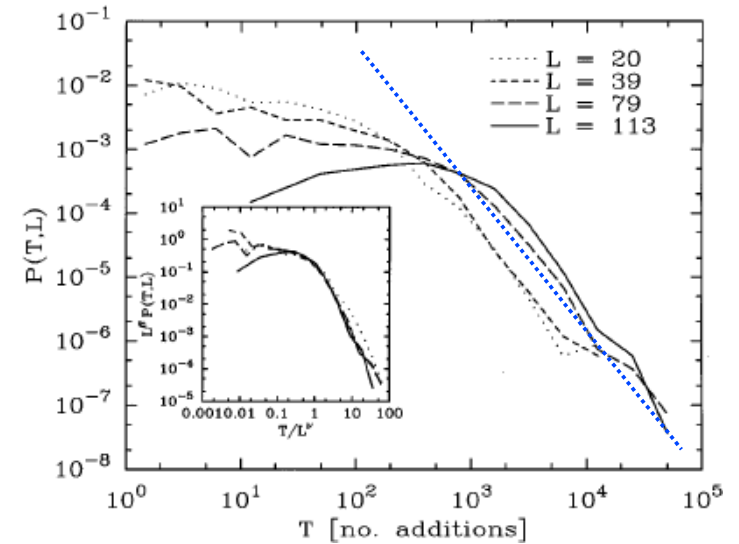
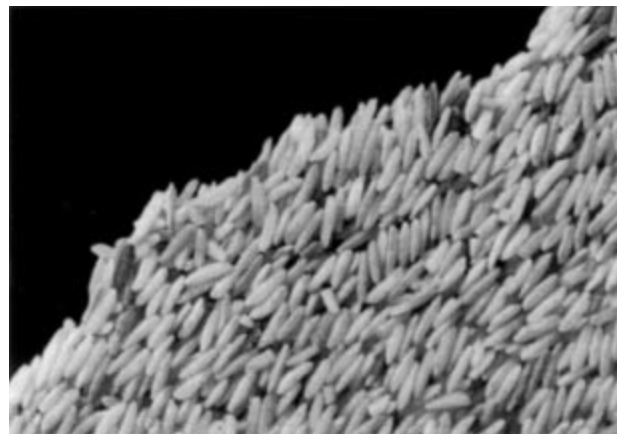
# But long-time dynamics governed by power-law waits



Waiting times  $\gg$  fluid timescales:  
 → Driven by burial and excavation of bed



Rice pile  
 [Christensen et al., PRL, 1996]  
 shows power-law residence times due to burial and excavation

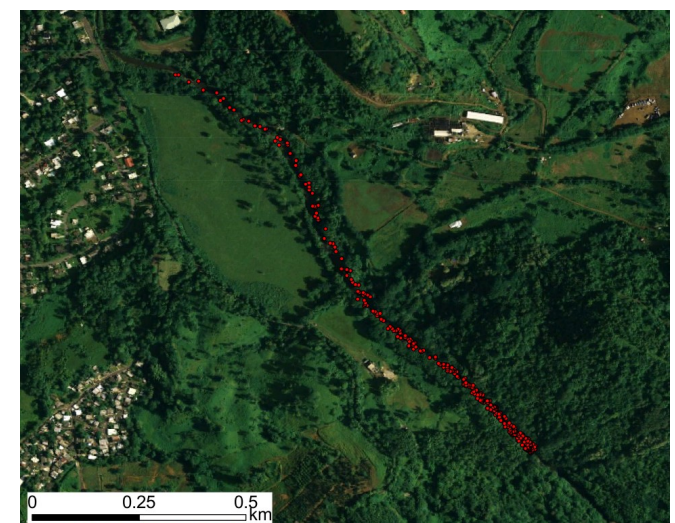
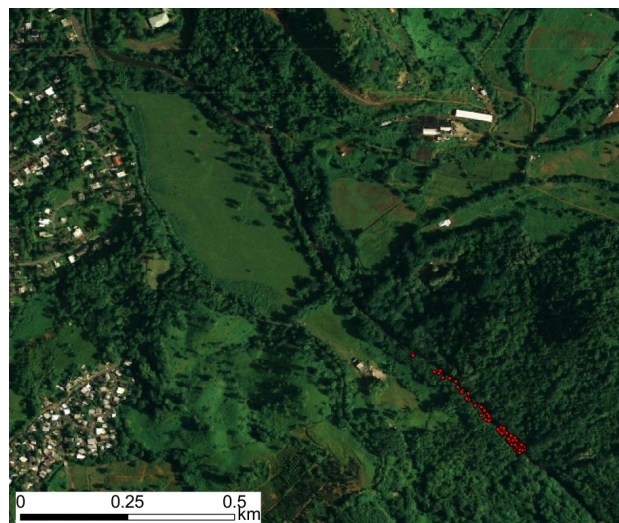
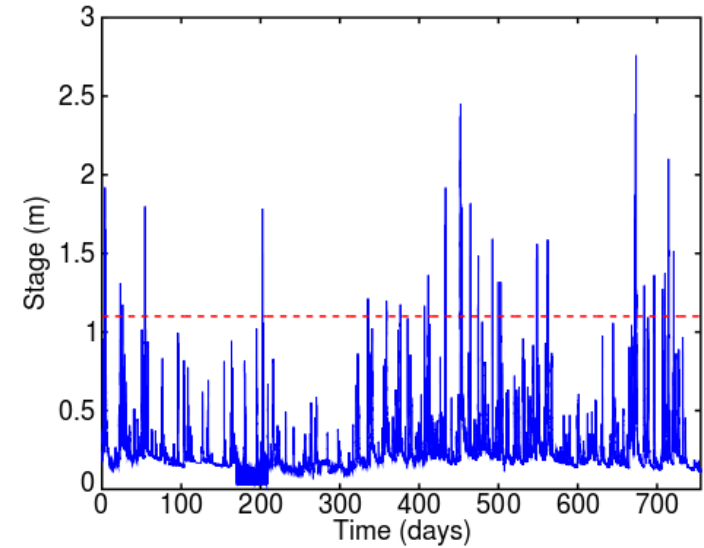




# Particle transport in real rivers: Radio Frequency Identifier (RFID) Tags

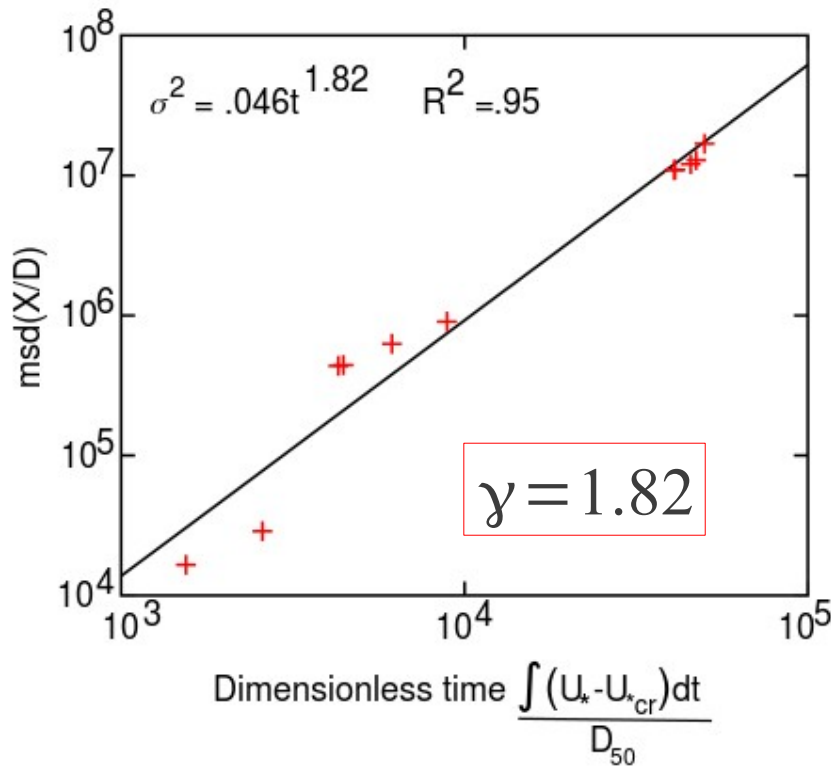
- Intermittent floods drive particle motion.

- Measure position of “radio rocks” after each flood.



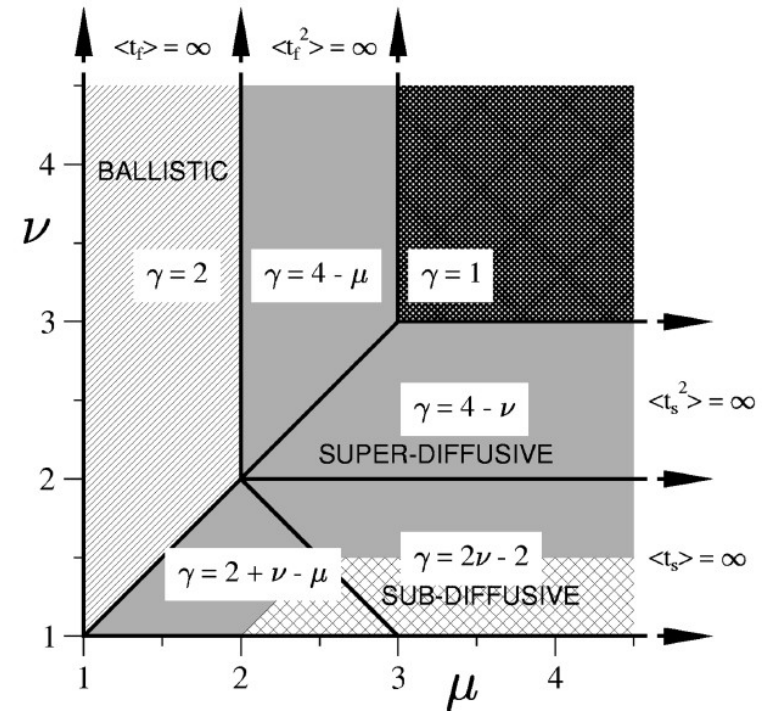
# Dispersion: Superdiffusion due to power-law waits + drift

Mean square displacement →  
Superdiffusion

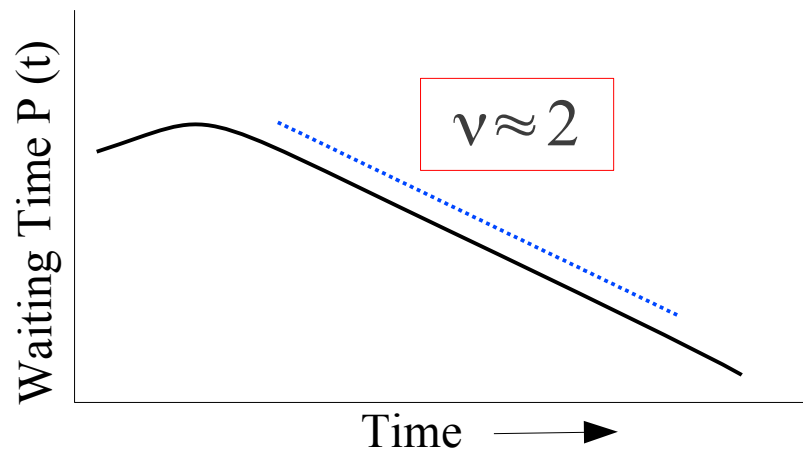


Anomalous diffusion resulting from strongly asymmetric random walks

Eric R. Weeks\* and Harry L. Swinney†



- Flights are thin tailed ( $\mu = \infty$ ).
- Transport is strongly asymmetric
- ∴ Power-law wait times.



Mobile particles →  
Momentum conservation.

Particles get stuck in bed.

Mobile ↔ immobile transitions w/ power-law waiting.

**Fractional advection-dispersion equations** for modeling transport at the Earth surface

Rina Schumer,<sup>1</sup> Mark M. Meerschaert,<sup>2</sup> and Boris Baeumer<sup>3</sup>

**Tracer particles spend much more time at rest than in motion.**

Stochastic modeling approach:

Direct solution of fADE, if known, to determine dispersion.

Lagrangian particle tracking to determine dispersion from collection of particle motions.

But how to assess, *a priori*, what particle waiting times and hop lengths are?

→ Need better understanding of physics

$$\frac{\partial^\alpha C(x,t)}{\partial t^\alpha} = -v \frac{\partial C(x,t)}{\partial x} + D \frac{\partial^2 C(x,t)}{\partial x^2}.$$

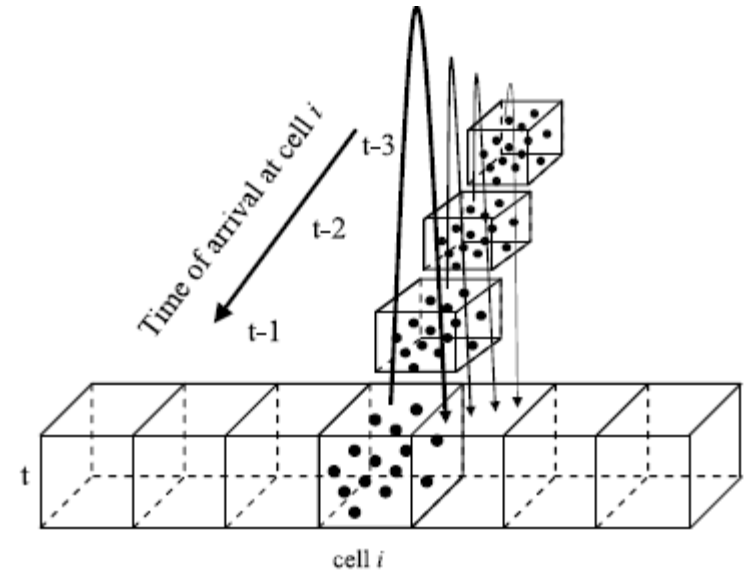
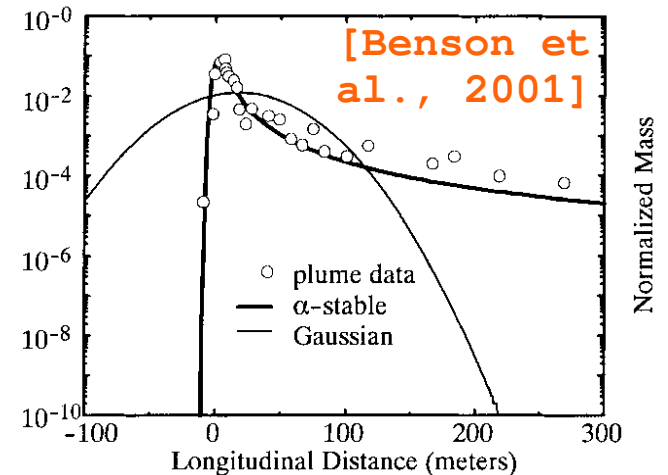


Figure 9. When governed by a fractional-in-time ADE, particles have memory of the time that they arrive at a given point. Their probability of release decays as a power law from arrival time.

Anomalous Dispersion

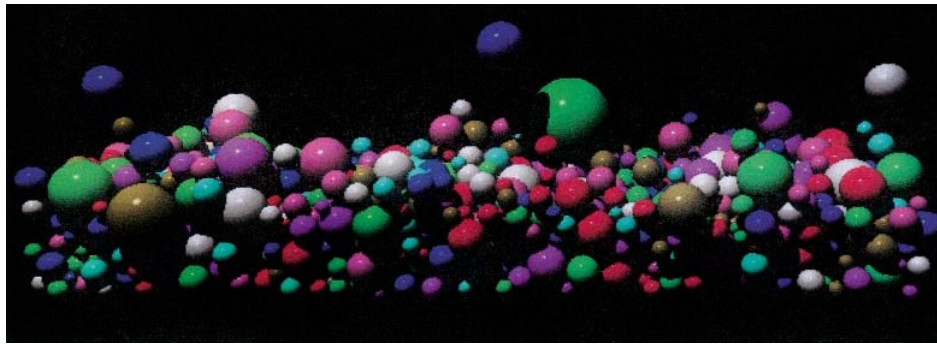


# Summary and directions

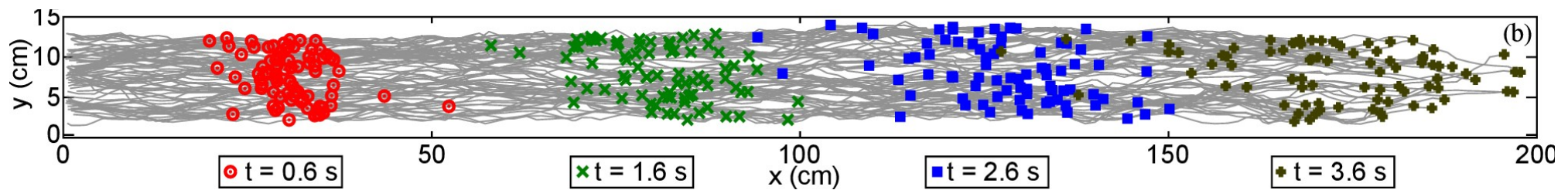
**Thresholds of motion:** stick-slip dynamics, stochastic transport



**Direct simulation:** possible path forward, difficult for natural systems



**Statistical mechanics:** useful framework for deriving transport equations but mobile/immobile partitioning complicates application



**Fractional ADEs and Random Walk models:** flexible for modeling anomalous diffusion, but must be informed by physics