

Maximum-likelihood Theory for the Inversion of Gravity and Topography

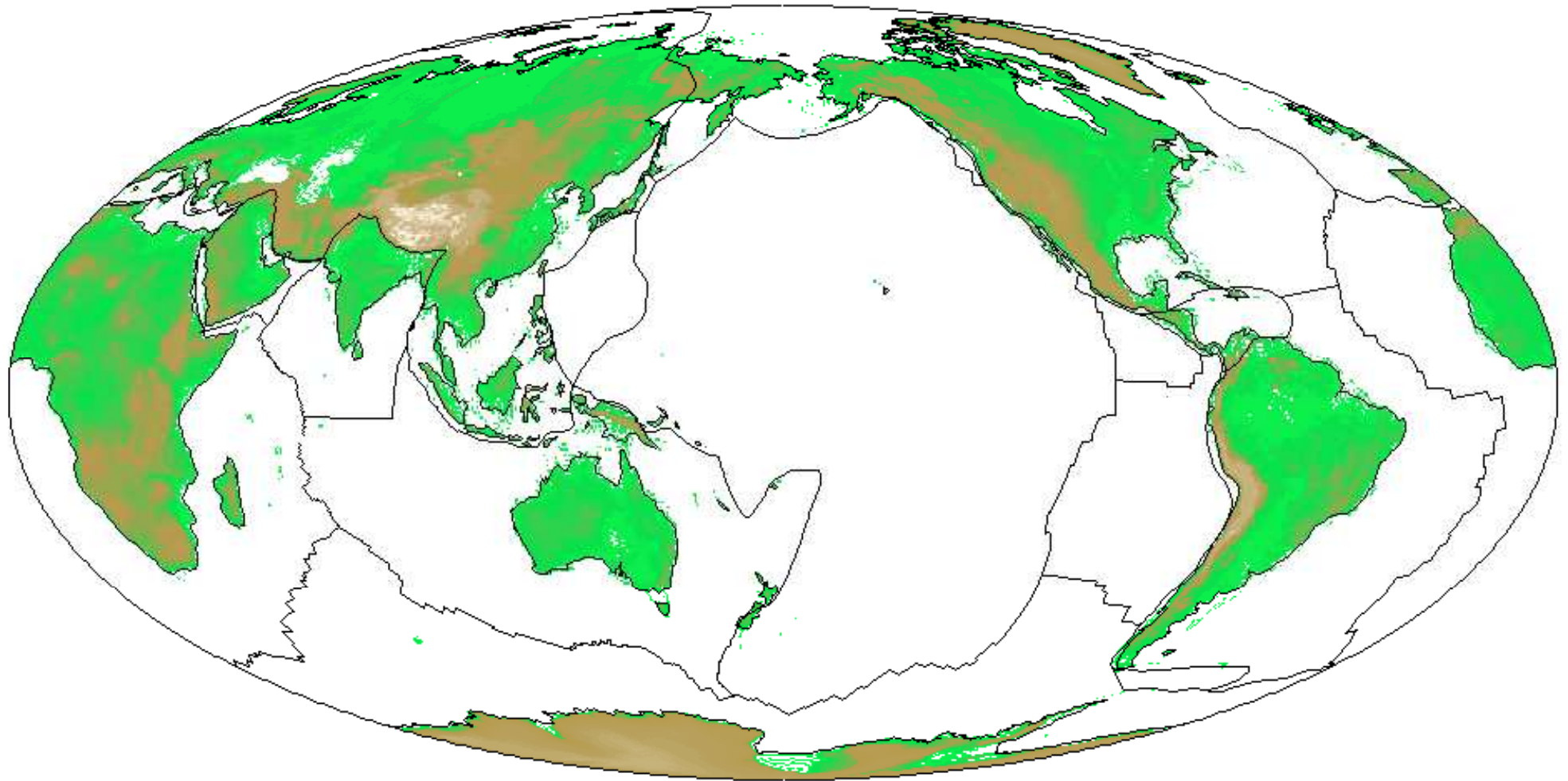
Frederik J Simons

Princeton University

Sofia C. Olhede

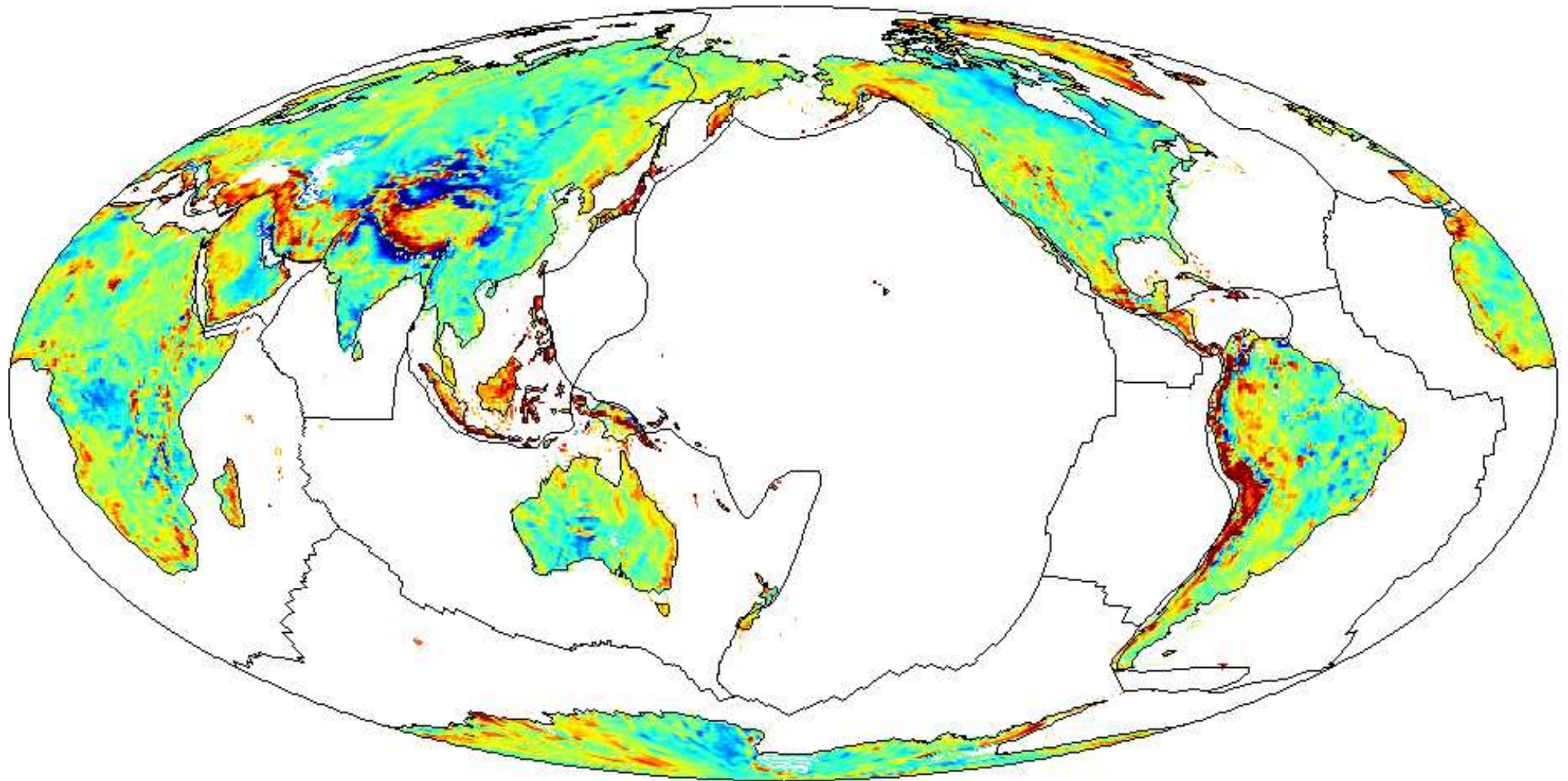
University College London





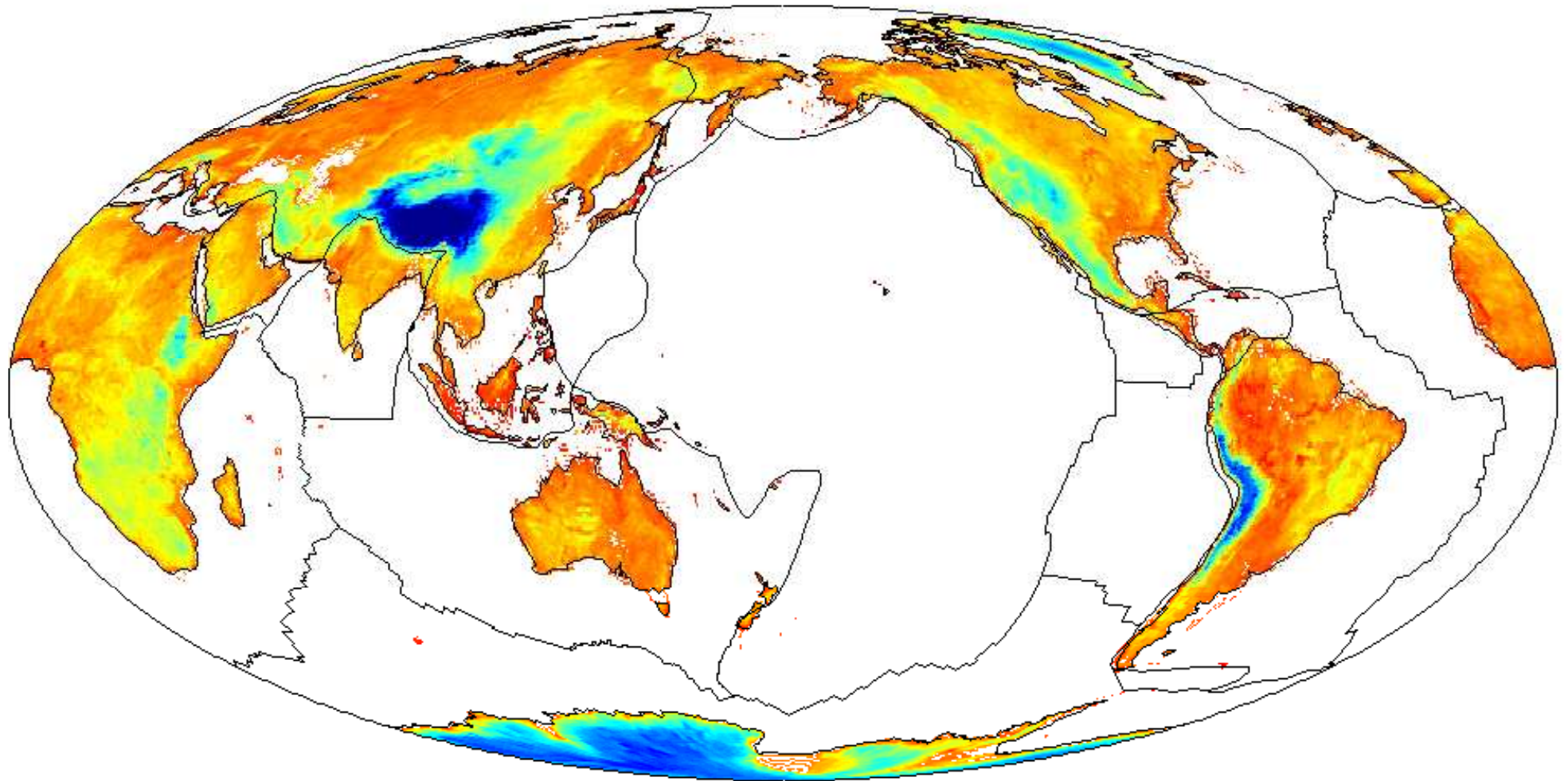
The free-air gravity anomaly

3/25

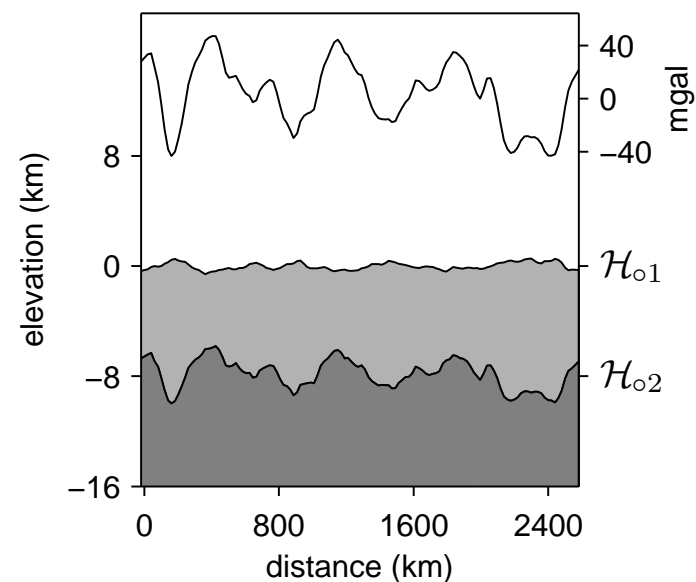
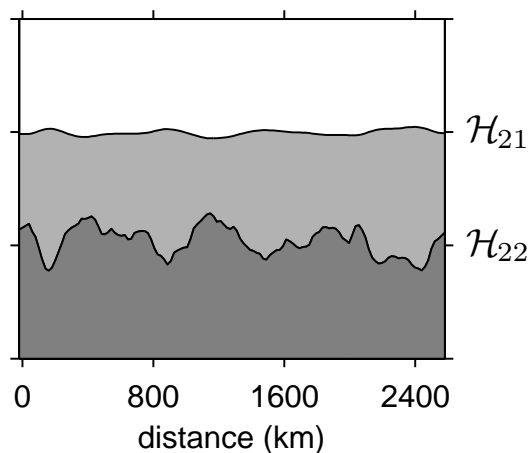
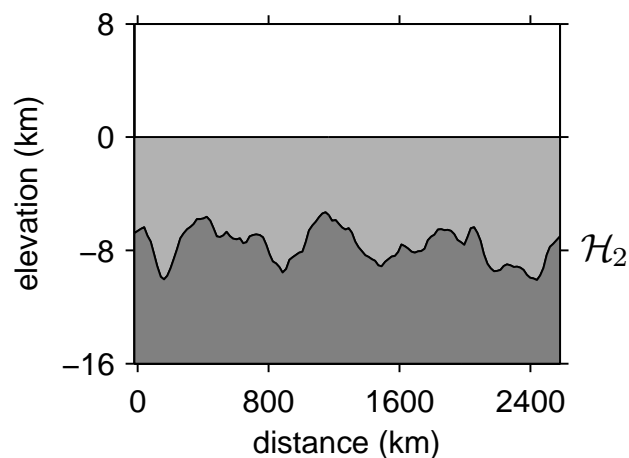
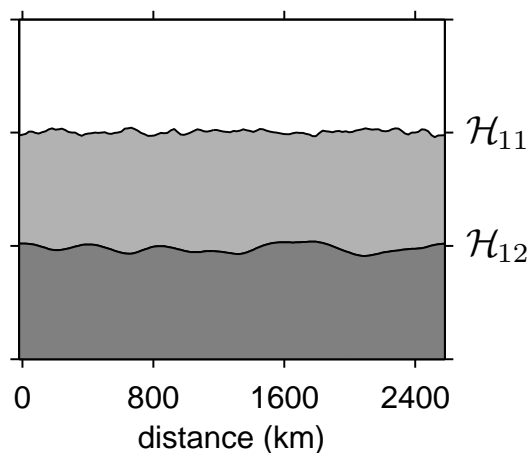
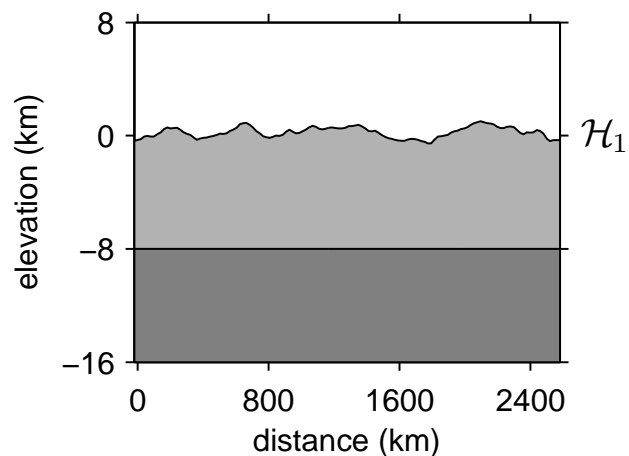


The Bouguer-air gravity anomaly

4/25



The standard model



$D = 7e+22 ; f^2 = 0.4$

$\sigma^2 = 0.0025$

$v = 2$

$\rho = 20000$

The lithosphere is two differential equations

Surface loading

Surface-loading topography \mathcal{H}_{11} is in instantaneous **elastic** balance with the sub-surface topography \mathcal{H}_{12} , according to the **biharmonic equation**:

$$\left(\nabla^4 + \frac{g\Delta_2}{D} \right) \mathcal{H}_{12}(\mathbf{x}) = -\frac{g\Delta_1}{D} \mathcal{H}_{11}(\mathbf{x}). \quad (1)$$

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Similarly, subsurface-loading topography \mathcal{H}_{22} is balanced at the surface by \mathcal{H}_{21} following the same equation:

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We want to find D , the **flexural rigidity**.

Thus the **free-air anomaly** $d\mathcal{G}_{ij}(\mathbf{k})$ due to the topographic perturbation $d\mathcal{H}_{ij}(\mathbf{k})$, at the j th interface resulting from the i th loading process, is given by

$$d\mathcal{G}_{ij}(\mathbf{k}) = 2\pi G \Delta_j d\mathcal{H}_{ij}(\mathbf{k}) e^{kz_j}. \quad (3)$$

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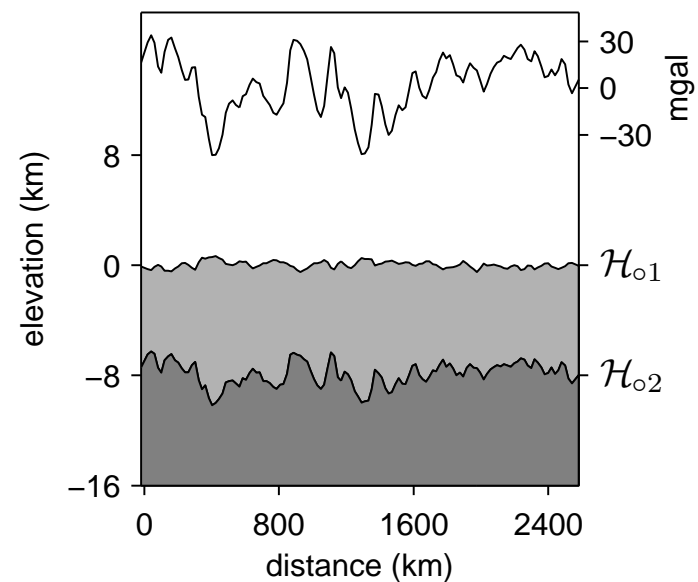
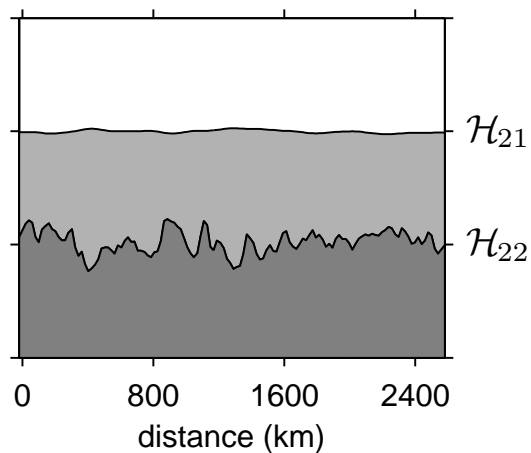
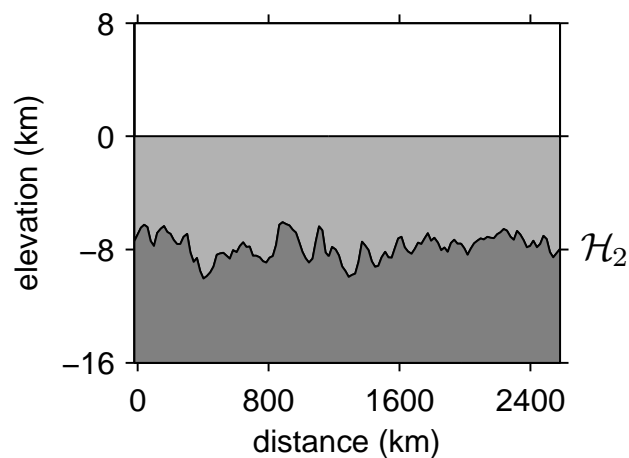
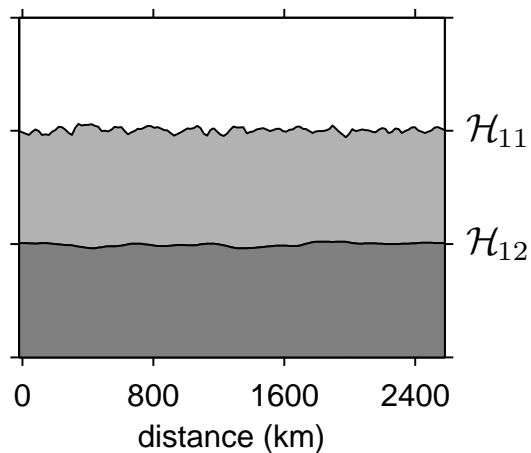
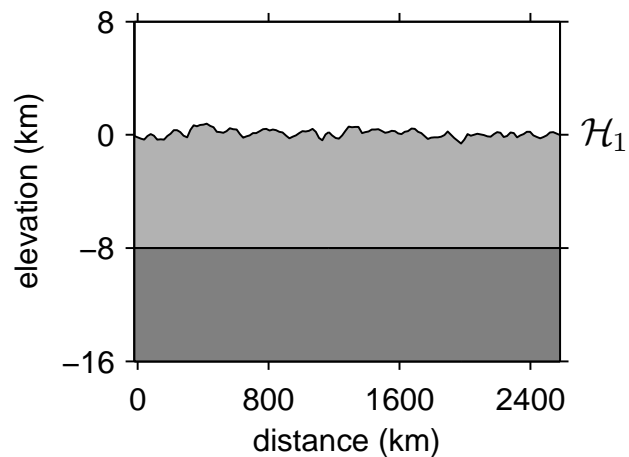
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The **Bouguer gravity anomaly** is calculated from the free-air anomaly by subtracting the gravitational effect from the observable surface topography $d\mathcal{H}_{\circ 1}(\mathbf{k})$,

$$d\mathcal{G}_{\circ 2}(\mathbf{k}) = d\mathcal{G}_{12}(\mathbf{k}) + d\mathcal{G}_{22}(\mathbf{k}). \quad (5)$$

The standard model — again



$D = 7e+22$; $f^2 = 0.4$; $r = -0.75$
 $\sigma^2 = 0.0125$
 $v = 2$
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What do we want to know?

9/25

- Given *a whole lot* of high-quality gravity and topography data, we want to find **one single parameter**, the rigidity, D , that describes their *relation*.

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- We allow ourselves to assume a certain **proportionality** between the power-spectral densities of the inputs: f^2 .
- We specify a joint structure to the initial inputs that also allows for their **correlation**: r .

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Define the **admittance**:

$$Q'_o(\mathbf{k}) = \frac{\langle d\mathcal{G}_{o2}(\mathbf{k}) d\mathcal{H}_{o1}^*(\mathbf{k}) \rangle}{\langle d\mathcal{H}_{o1}(\mathbf{k}) d\mathcal{H}_{o1}^*(\mathbf{k}) \rangle}.$$

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Define the **coherence-squared**:

$$\gamma_o'^2(\mathbf{k}) = \frac{|\langle d\mathcal{G}_{o2}(\mathbf{k}) d\mathcal{H}_{o1}^*(\mathbf{k}) \rangle|^2}{\langle d\mathcal{H}_{o1}(\mathbf{k}) d\mathcal{H}_{o1}^*(\mathbf{k}) \rangle \langle d\mathcal{G}_{o2}(\mathbf{k}) d\mathcal{G}_{o2}^*(\mathbf{k}) \rangle}.$$

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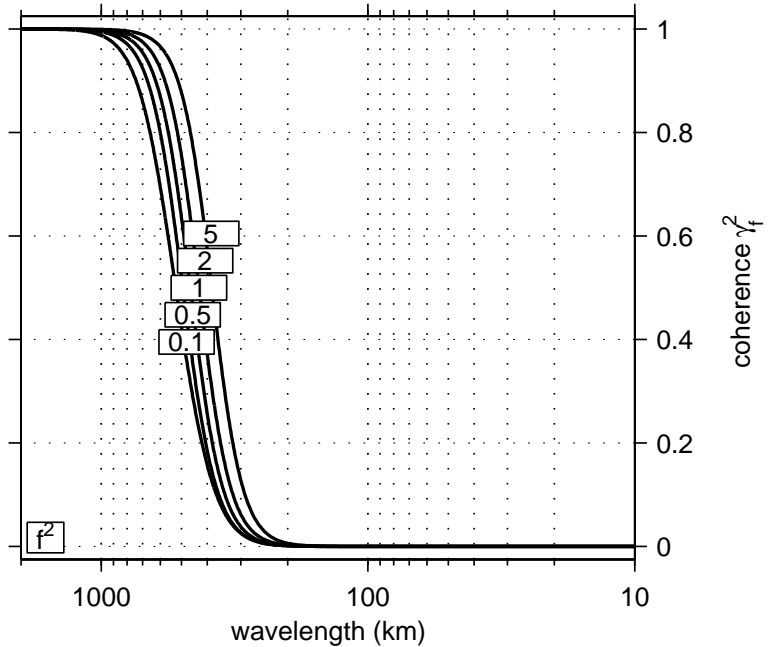
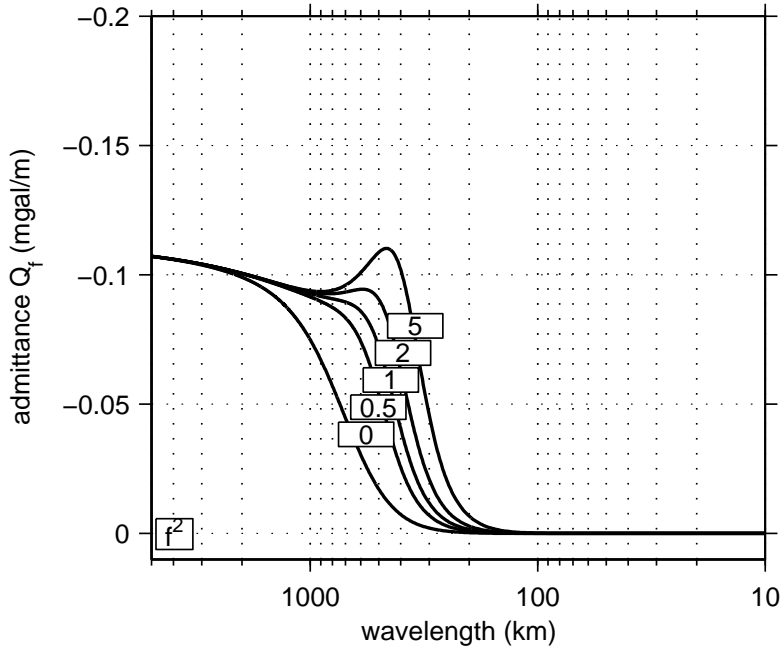
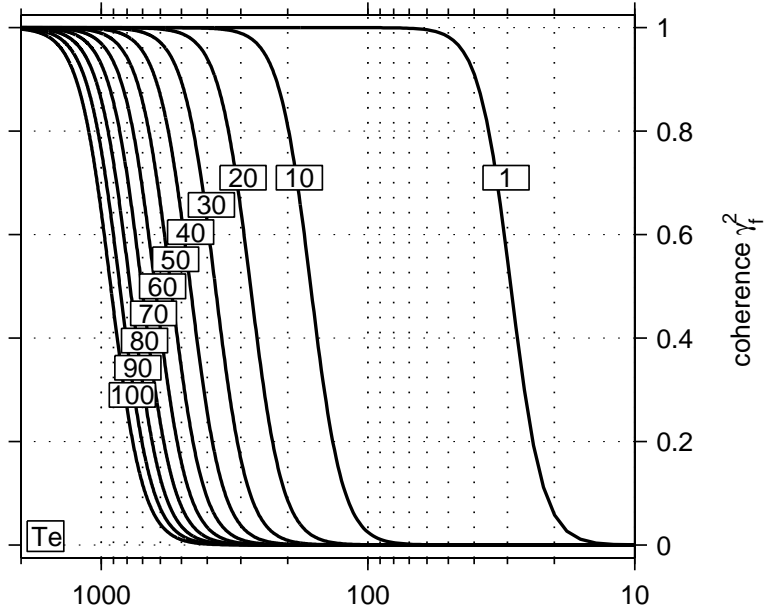
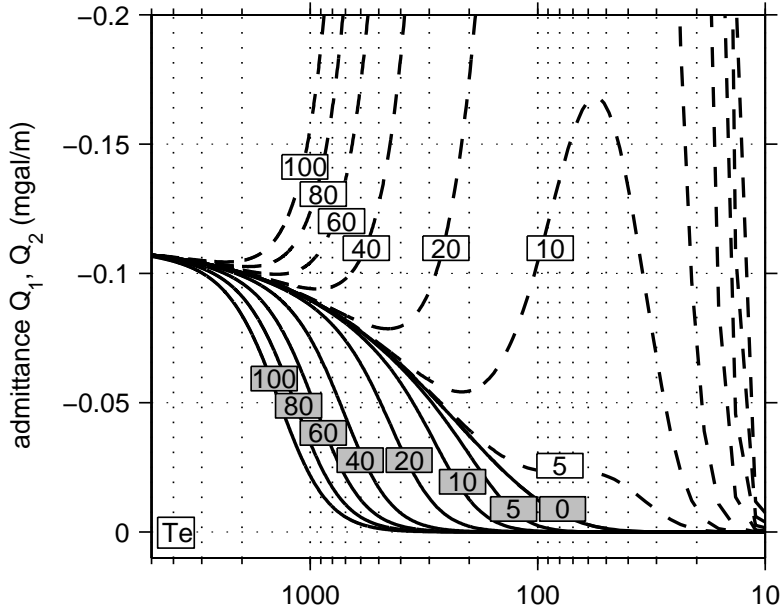
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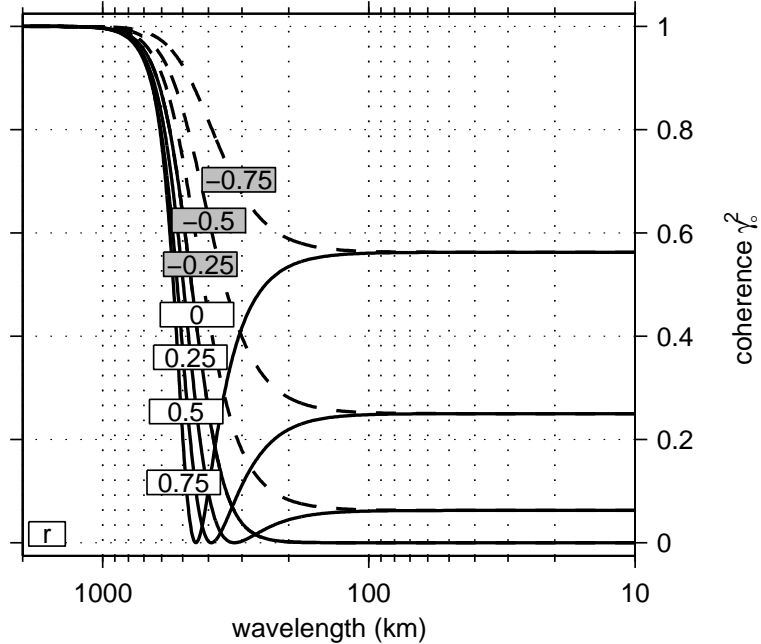
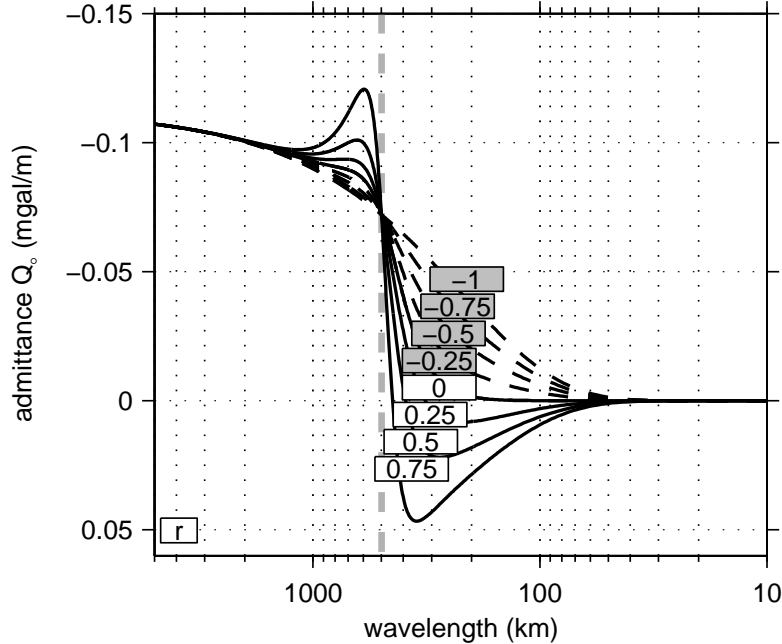
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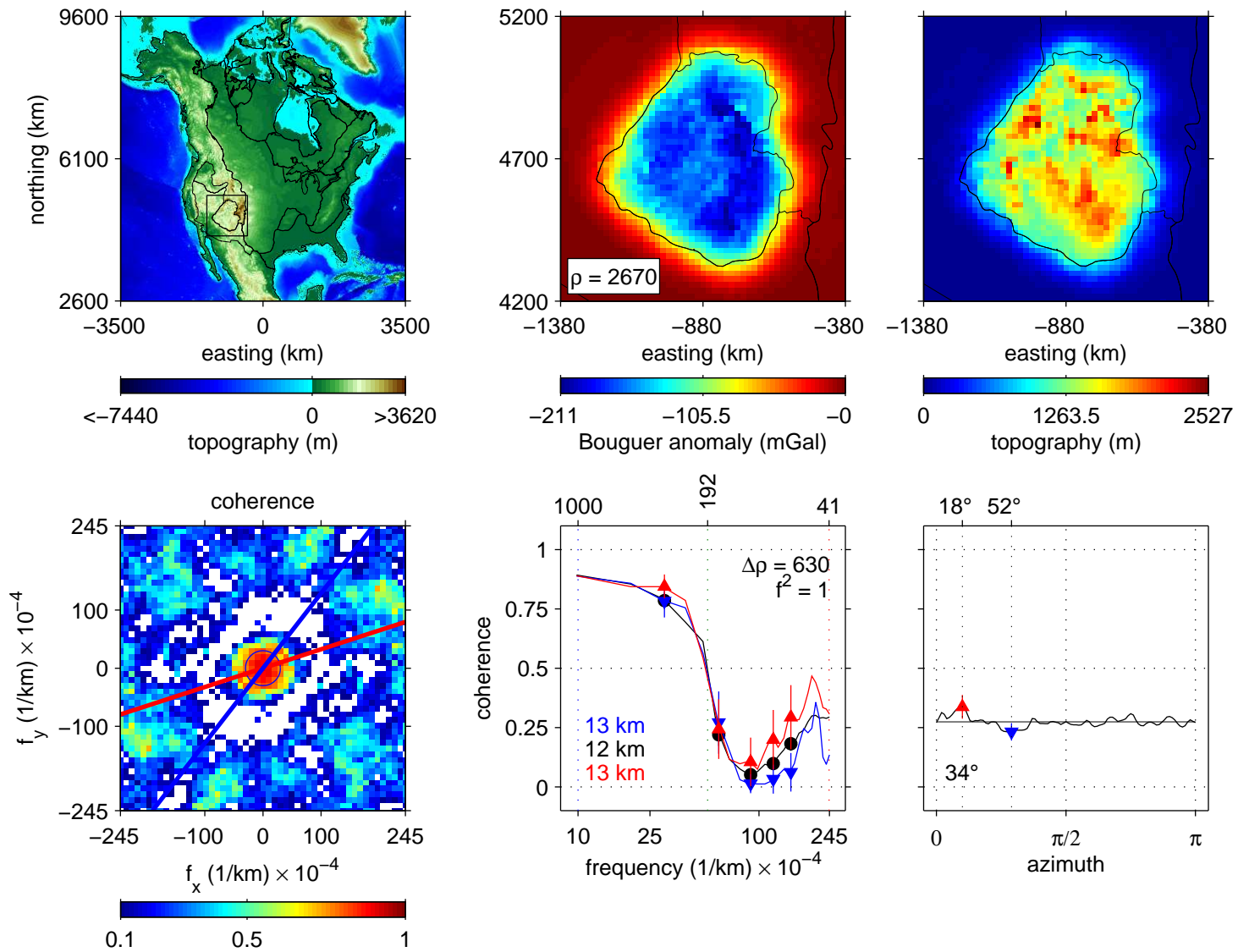
The forward model is a very doable function of D , f^2 , and r .

Admittance and Coherence — I

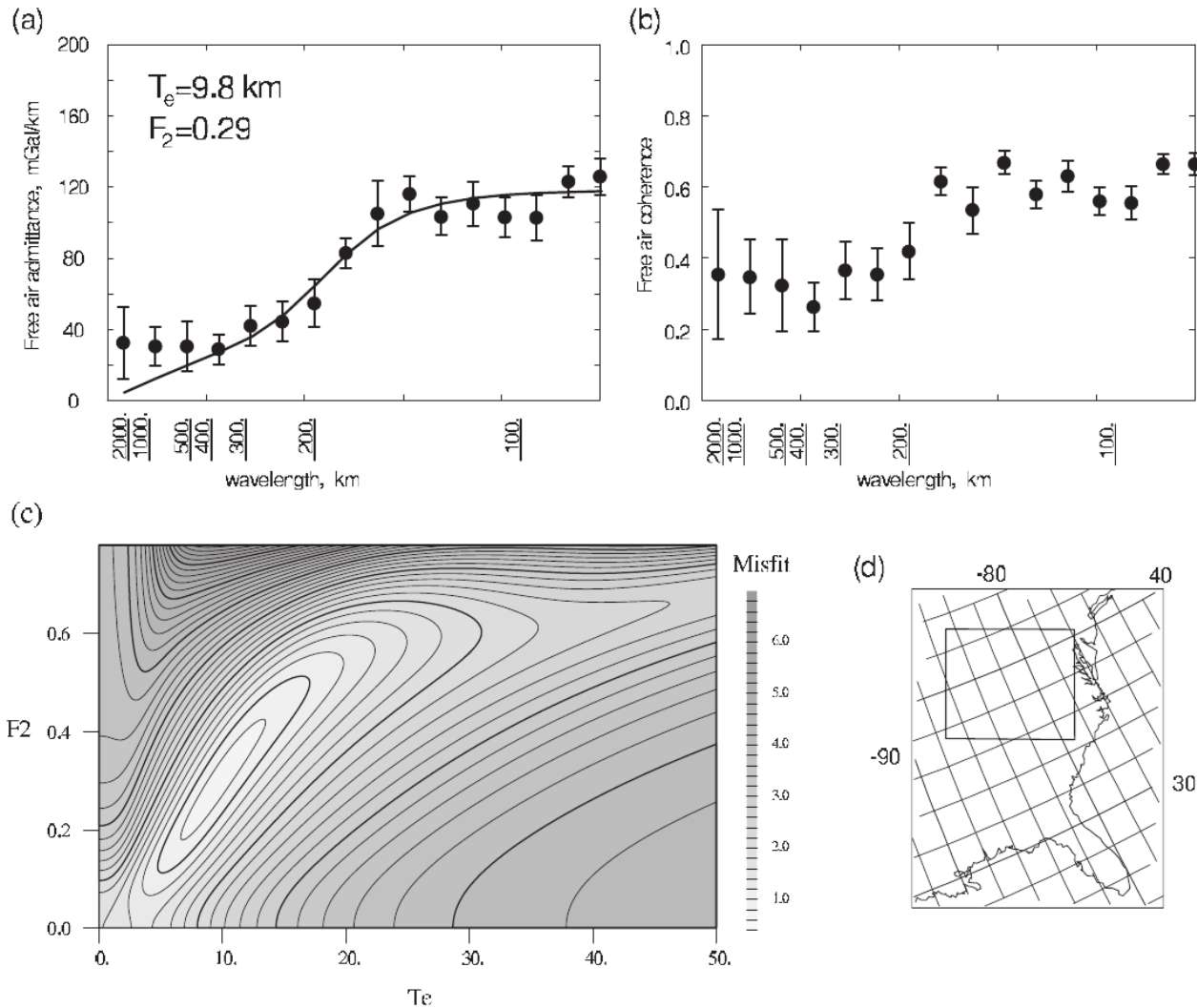


Admittance and Coherence — III





eastern US



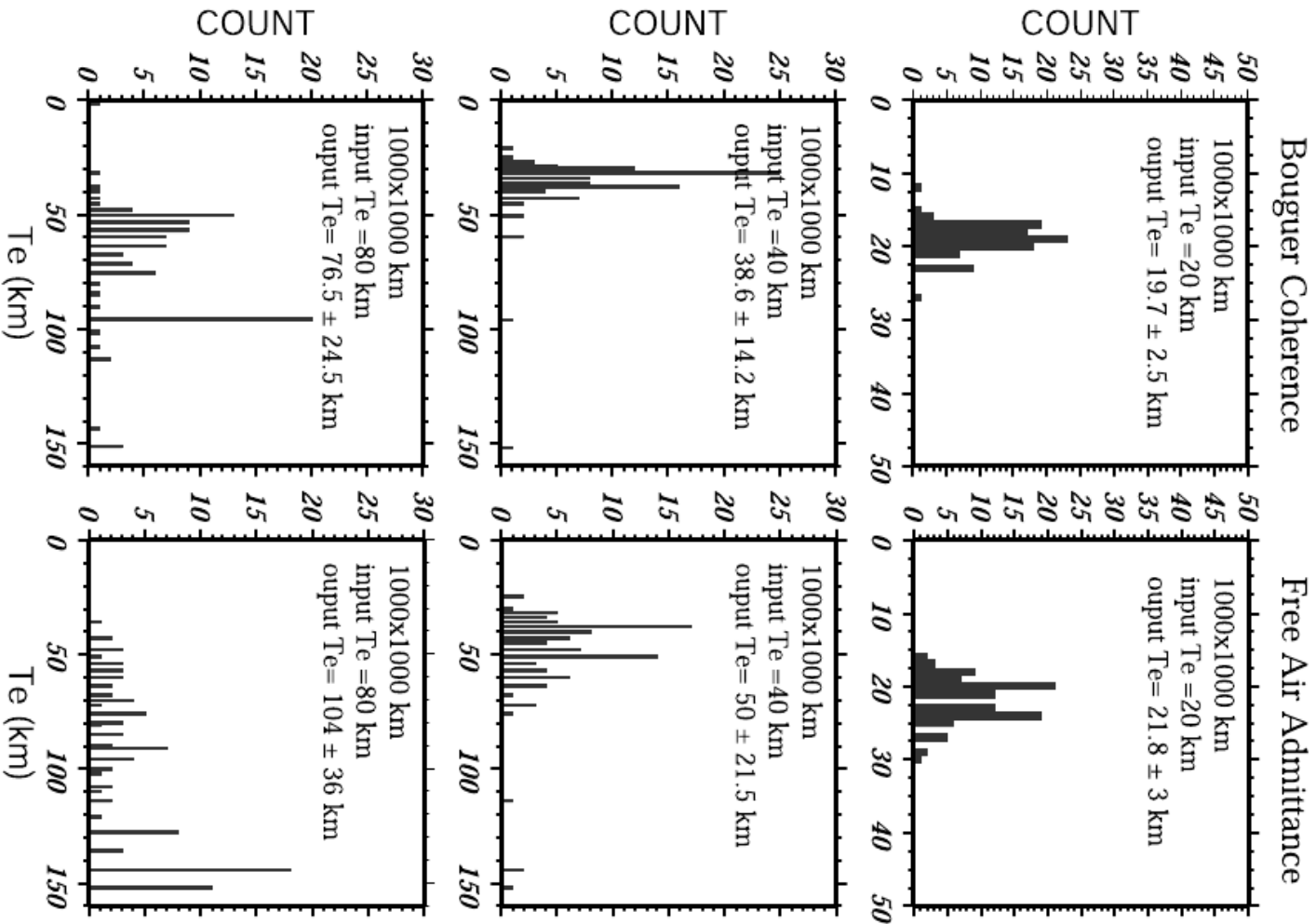


Figure 2. (left) Histograms of 100 Bouguer coherence and (right) free-air admittance analyses of synthetic data generated with input or “true” $T_e = 20, 40,$ and 80 km and an average loading ratio of $f=1$. The window size used for analysis was 1000×1000 km. Output T_e is the mean and standard deviation resulting from the 100 analyses.

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16/25

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18/25

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Note that the forms for admittance and coherence are ratios of the terms in this variance matrix. In our formulation, we simply retain the individual terms, which lead to Gaussian forms, rather than forming the spectral ratios, which don't.

A direct maximum-likelihood solver for D, f^2, r 19/25

Now we can form the **log-likelihood** of observing \mathbf{H} under their being given by a loading process S_{11} and a lithospheric response characterized by D and f^2 :

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- Assuming **isotropy**, we choose a three-parameter **Matérn** form for the initial driving loads. **Anisotropy** is next.

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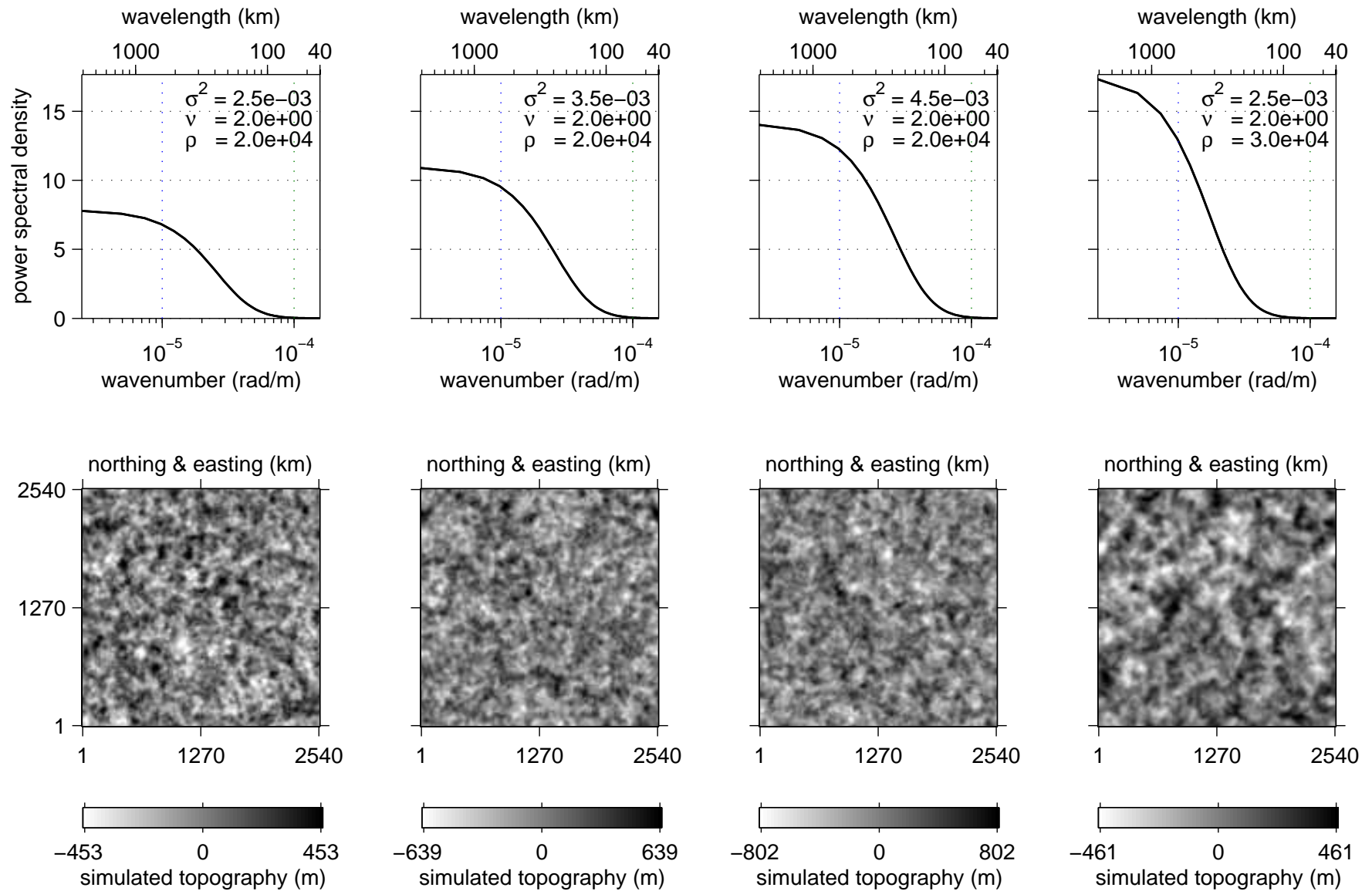
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The **maximum-likelihood** is the **best**, minimum-variance, unbiased estimate of the new (three lithospheric, three spectral) parameter vector:

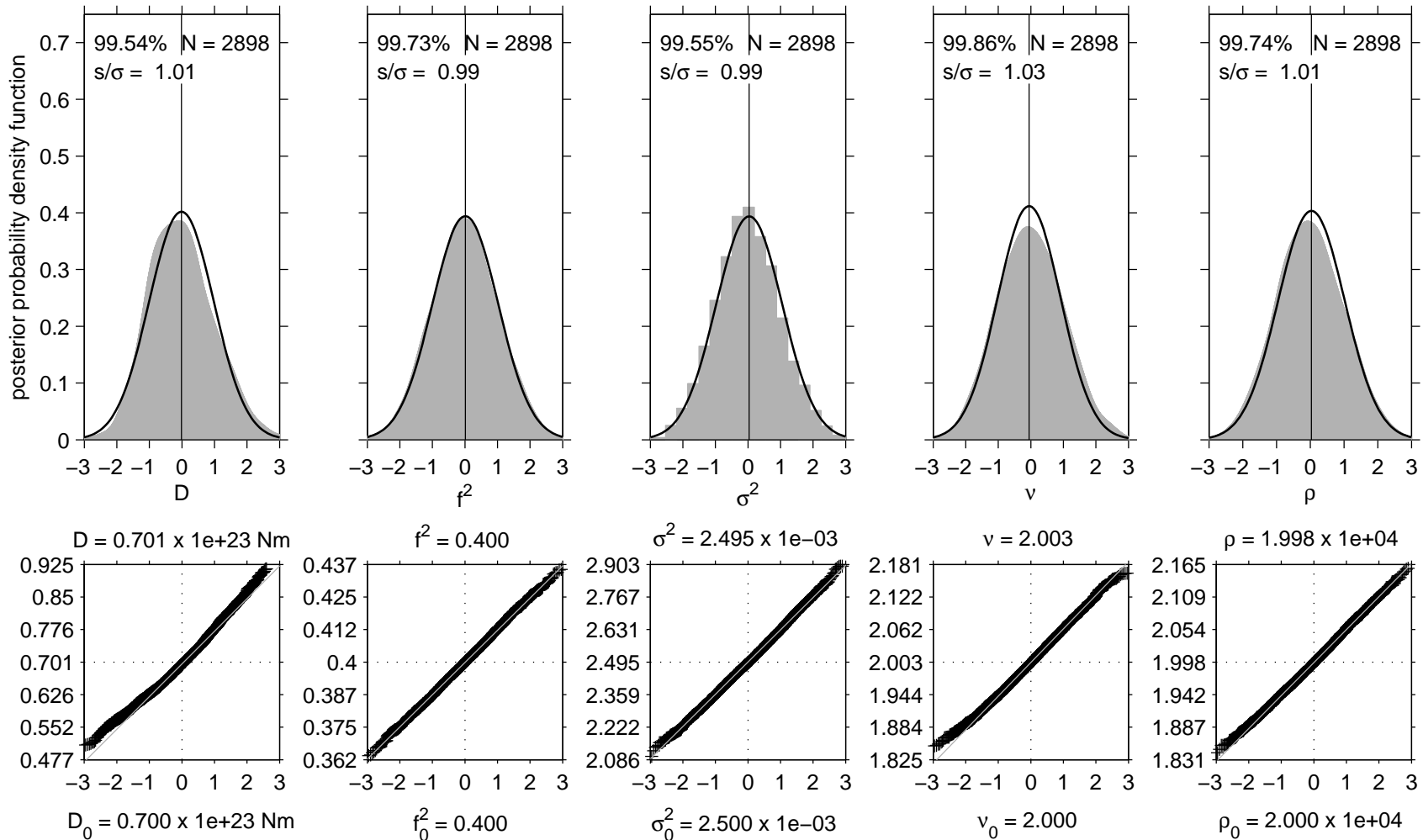
$$\boldsymbol{\theta} = [D \ f^2 \ r \ \sigma^2 \ \nu \ \rho]^T.$$

The (isotropic) Matérn spectral class



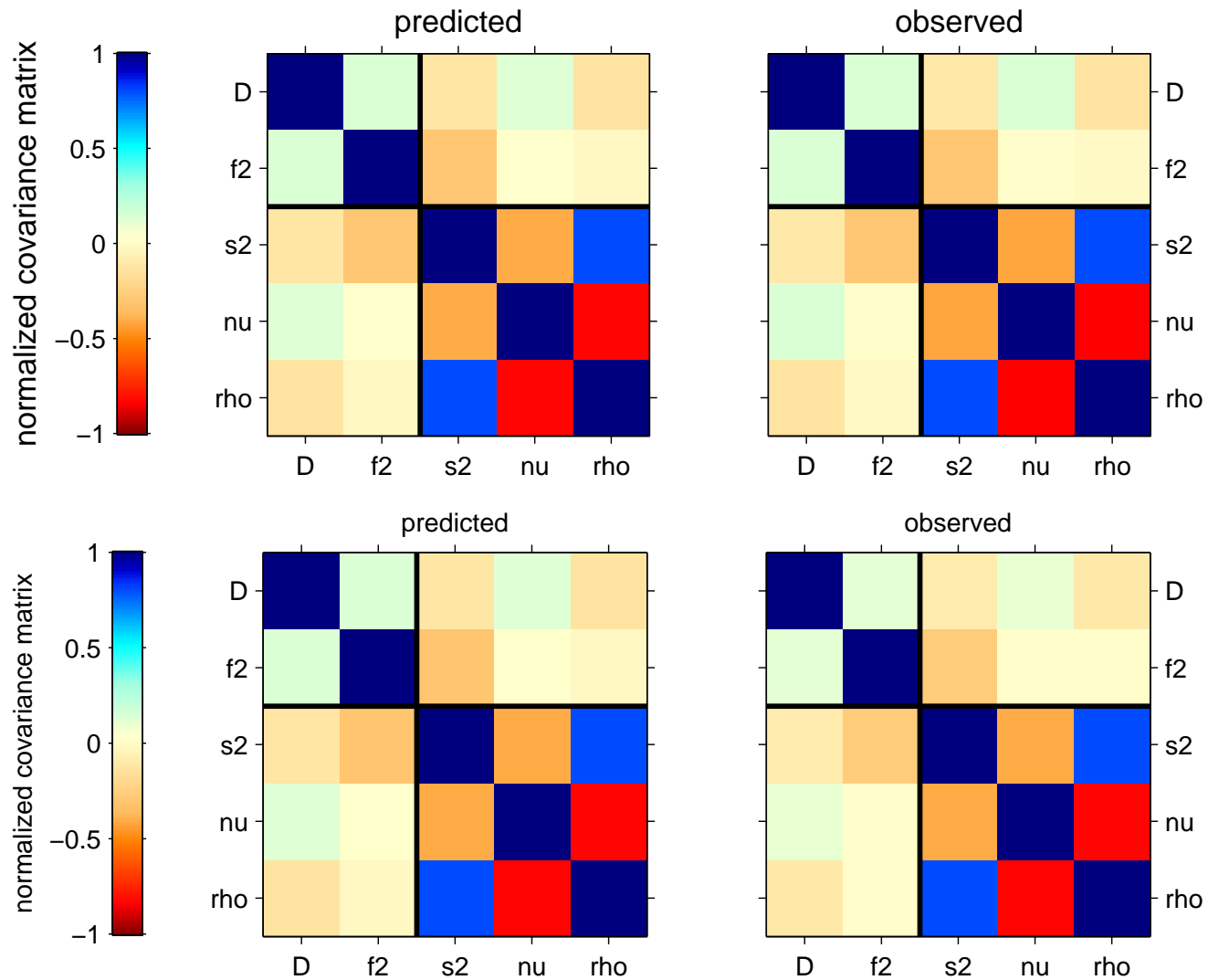
Does it work? Results

MLE Simulations with $\Delta_1 = 2670$; $\Delta_2 = 630 \text{ kg m}^{-3}$; $z_2 = 35 \text{ km}$; 64×64 grid ; $1280 \times 1280 \text{ km}$

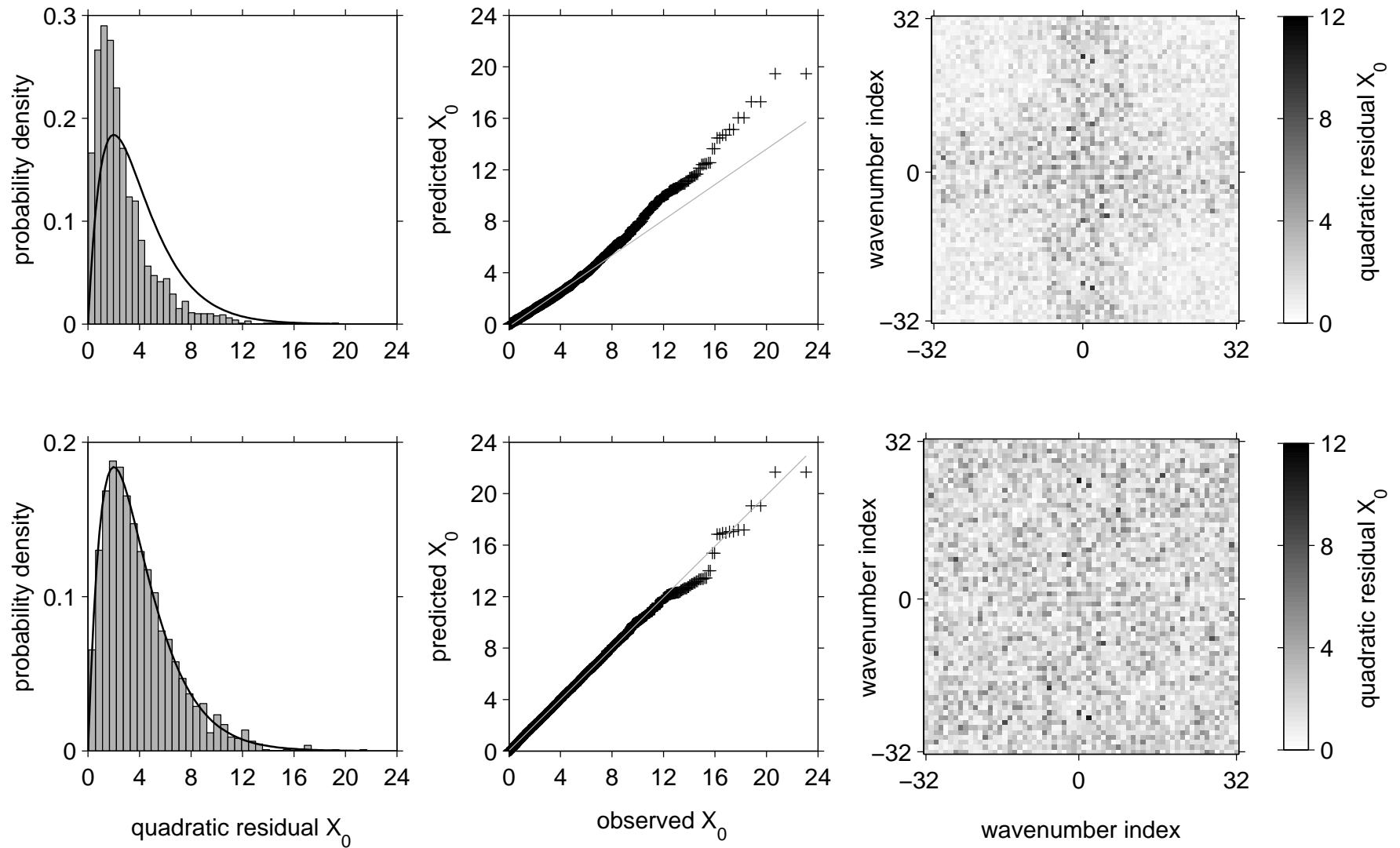


With $E = 1.4 \times 10^{11}$; $\nu = 0.25$; $T_e = 18 \pm 8 \text{ km}$; $f^2 = 0.4 \pm 0.013$

Does it work? Trade-offs



Does it work? Modeling the residuals



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- There are issues with convincing some very prominent geoscientists that they need *this much* statistics.

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- The connections in marrying **deterministic** forward models with **stochastic** inputs and observables are fruitful and widespread.

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- Simulations verify that the estimates are normally distributed, and unbiased and closely track the variance predicted by the **fully analytic theory**, which is minimized. Confidence intervals are symmetric, and covariance between the estimated D and f^2 is small. Though no longer needed, predicting admittance and coherence in retrospect vastly tightens their error bars.

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 - Our current method successfully takes into account r -**correlated** loads, small data sets, irregular domains. **Anisotropy** in both the loading and the response is further down the line. *Then*, the statistics will be interesting.
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