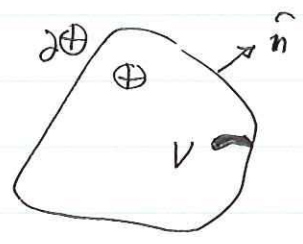


Class #9

Stress glut  $\underline{\underline{S}} = \underline{\underline{T}}^{Hooke} - \underline{\underline{T}}^{true}$

Equivalent body force  $\underline{\underline{f}} = -\nabla \cdot \underline{\underline{S}}$  in  $\oplus$

Equivalent surface force  $\underline{\underline{t}} = \hat{n} \cdot \underline{\underline{T}}$  on  $\partial\oplus$



source volume  $V$

Total force  $\underline{\underline{F}}$  and torque  $\underline{\underline{N}}$  are zero (latter because  $\underline{\underline{S}} = \underline{\underline{S}}^T$ )

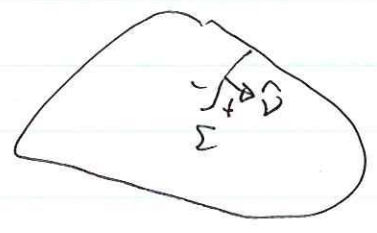
Acceleration response

$$\underline{\underline{a}}(\underline{\underline{x}}, t) = \sum_R (a_k \cos \omega_k t + b_k \sin \omega_k t) \underline{\underline{S}}_k(\underline{\underline{x}})$$

$$\begin{Bmatrix} a_k \\ b_k \end{Bmatrix} = \int_{t_0}^{t_f} \int_V \frac{\partial \underline{\underline{S}}}{\partial t} : \underline{\underline{\epsilon}}_k \begin{Bmatrix} \cos \omega_k t \\ \sin \omega_k t \end{Bmatrix} dV dt$$

integrate only over source volume  $V$  and interval  $t_0 \leq t \leq t_f$

Ideal fault



tang. slip  $\underline{\underline{\Delta S}}(\underline{\underline{x}}, t) = \underline{\underline{S}}(\underline{\underline{x}}^+, t) - \underline{\underline{S}}(\underline{\underline{x}}^-, t)$   
 $\hat{v} \cdot \underline{\underline{\Delta S}} = 0$

moment tensor density :  $m_{ij} = C_{ijkl} v_j \Delta S_k$   
 $\underline{\underline{m}} = \underline{\underline{C}} : \hat{v} \underline{\underline{\Delta S}}$

$$\frac{\text{moment}}{\text{area}} = \frac{m \omega r}{\text{time}^2} \quad \text{units}$$

Then  $\underline{\underline{S}} = \underline{\underline{m}} \delta_{\Sigma}$ , confined to fault surface, depends only on slip  $\underline{\underline{S}}$  and orientation  $\hat{n}$

$[\underline{\underline{C}}]_{\pm}$  across  $\Sigma$

Equivalent body force densities

$$\underline{\underline{f}} = - \underline{\underline{m}} \cdot \nabla \delta_{\Sigma} \quad \text{in } \oplus$$

$$\underline{\underline{t}} = (\hat{n} \cdot \underline{\underline{m}}) \delta_{\Sigma} \quad \text{on } \partial \oplus$$

Completely general: dynamic time-dependent faulting in an anisotropic  $\oplus$

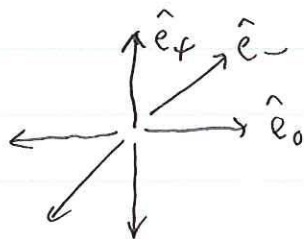
Symmetry  $\underline{\underline{m}} = \underline{\underline{m}}^T \Rightarrow$  diagonalizable

$$\underline{\underline{m}} = m_+ \hat{e}_+ \hat{e}_+ + m_0 \hat{e}_0 \hat{e}_0 + m_- \hat{e}_- \hat{e}_-$$

Equivalent body force density

$$f_j = - m_{ij} \partial_j \delta_{\Sigma}$$

Three linear vector dipoles



$m_{+,0,-} > 0$  : outward  
 $m_{+,0,-} < 0$  : inward

no contr. since  $v_k \Delta s_k = 0$

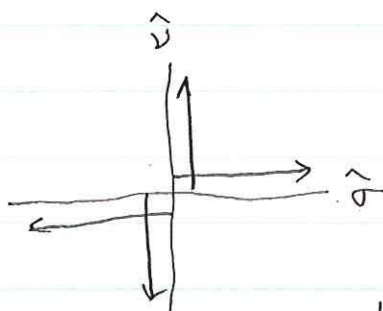
Isotropic  $\oplus$  :  $c_{ijkl} = (\kappa - \frac{2}{3}) \delta_{ij} \delta_{kl} + \mu (\delta_{ik} \delta_{jl} + \delta_{il} \delta_{jk})$

$$m_{ij} = \mu \Delta s (v_i \sigma_j + \sigma_i v_j)$$

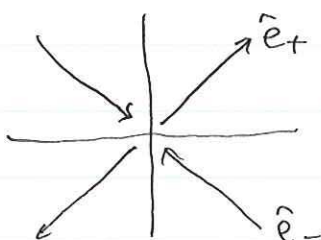
where  $\underline{\Delta s} = \Delta s \hat{\sigma}$

$$\underline{\underline{m}} = \mu \Delta s (\hat{\sigma} \hat{\sigma} + \hat{\sigma} \hat{\sigma})$$

Equivalent body force a double couple



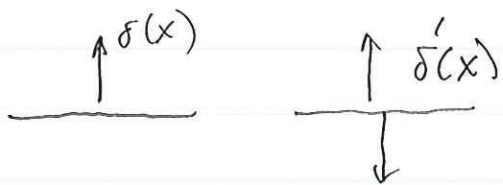
or



obviously no net force or torque

~~$$\underline{\underline{m}} = \mu \Delta s (\hat{e}_+ \hat{e}_+ - \hat{e}_- \hat{e}_-)$$~~

$$\hat{e}_{\pm} = \frac{1}{\sqrt{2}} (\hat{\sigma} \pm \hat{\tau})$$



Valid for a curved ~~fault~~ or corrugated fault

Mode response to faulting

$$\begin{Bmatrix} a_k \\ b_k \end{Bmatrix} = \int_{t_0}^{t_f} \int_{\Sigma} \frac{\partial \underline{m}}{\partial t} : \underline{\underline{\Sigma}}_k \begin{Bmatrix} \cos \omega_k t \\ \sin \omega_k t \end{Bmatrix} dV dt$$

integrate only ~~the~~ over portion of  $\Sigma$  where slip is actively occurring.

Point source approximation : moment tensor

long wavelengths  $\Rightarrow$  source dimensions  
 long period  $\Rightarrow$  source duration

$$\underline{\underline{\Sigma}}_k(\underline{x}) \begin{Bmatrix} \cos \omega_k t \\ \sin \omega_k t \end{Bmatrix} \approx \underline{\underline{\Sigma}}_k(\underline{x}_s) \begin{Bmatrix} \cos \omega_k t_s \\ \sin \omega_k t_s \end{Bmatrix}$$

$\underline{x}_s, t_s$  hypocenter location in point source approximation

$$\begin{Bmatrix} a_k \\ b_k \end{Bmatrix} = \underline{\underline{M}} : \underline{\underline{\Sigma}}_k(\underline{x}_s) \begin{Bmatrix} \cos \omega_k t_s \\ \sin \omega_k t_s \end{Bmatrix}$$

where  $\underline{\underline{M}} = \int_{t_0}^{t_f} \int_V \frac{\partial \underline{S}}{\partial t} dV dt$ , moment tensor

$$= \int_V \underline{\underline{S}}^{final state} dV$$

volume integral of the final static stress  $\underline{\underline{S}}$ .

$$\cos a \cos b + \sin a \sin b = \cos(a-b)$$

5

Response

$$\underline{a}(\underline{x}, t) = \sum_k \underbrace{\underline{M} : \underline{\underline{\epsilon}}_k(\underline{x}_s)}_{\text{ampl.}} \underbrace{\underline{s}_k(\underline{x})}_{\text{shape}} \underbrace{\cos \omega_k (t - t_s)}_{\text{begins at } t=t_s \text{ origin time}}$$

Linear in moment tensor  $\underline{M}$   
Read quote from Gilbert

Moment tensor for an <sup>ideal</sup> earthquake fault

$$\underline{M} = \int_{\Sigma} \underline{m}^{\text{static final}} dA \quad \left( \int_{\oplus} \underline{m}^{\text{final}} \underline{\delta}_{\Sigma} dV \right)$$

Isotropic  $\oplus$  :

$$\underline{M} = \int_{\Sigma} \mu \Delta s^{\text{final}} (\hat{\nu} \hat{\nu} + \hat{\sigma} \hat{\sigma}) dA$$

Glut rate in point source approx

$$\underline{\dot{s}} = \underline{M} \delta(\underline{x} - \underline{x}_s) \delta(t - t_s)$$

Body force  $\underline{f} = -M \cdot \nabla \delta(\underline{x} - \underline{x}_s) \delta(t - t_s)$

3 linear vector dipoles

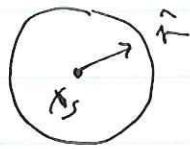
Planar fault, uni-directional slip:

$$\underline{M} = \text{classic double couple} \quad M_0 (\hat{\nu} \hat{\nu} + \hat{\sigma} \hat{\sigma}) \quad \text{moment} \quad M_0 = \int_{\Sigma} \mu \Delta s^{\text{final}} dA \quad \text{Aki (1966)}$$



skip this

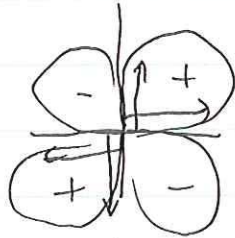
P-wave radiation pattern:



focal sphere

$$\hat{r} \cdot \underline{M} \cdot \hat{r}$$

beach ball



quadrupole pattern



talk about later

CMT = centroid - moment tensor

Solve for  $\underline{M}$  as well as an updated location  $\underline{x}_s + \underline{\Delta x}$ ,  $t_s + \Delta t$

$$\begin{aligned} \underline{a}(\underline{x}, t) = & \sum_k \underline{M} : \underline{\underline{\epsilon}}_k(\underline{x}_s) \underline{\underline{\epsilon}}_k(\underline{x}) \cos \omega_k (t - t_s) \\ & + \sum_k \underline{\Delta x} \underline{M} : \nabla \underline{\underline{\epsilon}}_k(\underline{x}_s) \underline{\underline{\epsilon}}_k(\underline{x}) \cos \omega_k (t - t_s) \\ & + \sum_k \omega_k \Delta t \underline{M} : \underline{\underline{\epsilon}}_k(\underline{x}_s) \underline{\underline{\epsilon}}_k(\underline{x}) \sin \omega_k (t - t_s) \end{aligned}$$

Linear inverse problem for  $\underline{M}$ ,  $\underline{\Delta x}$ ,  $\Delta t$

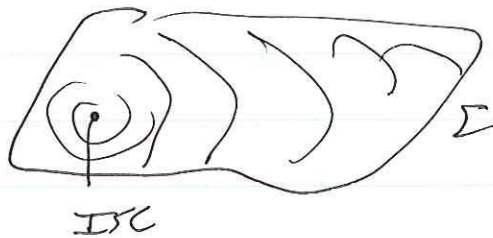
Can be shown (see D&T 5.4.2) that for a planar fault, uni-directional dip

$$\underline{\Delta x} = \frac{1}{M_0} \int_{\Sigma} (\underline{x} - \underline{x}_s) \mu \Delta s^f dA$$

↖ on fault

$$\Delta t = \frac{1}{M_0} \int_{t_0}^{t_f} \int_{\Sigma} (t - t_s) \mu d_t \Delta s dA dt$$

Centroid may differ from ISC location =  
beginning of rupture



Number of unknowns in a CMT inversion: 10

### 5.4.3 Deviatoric & double-couple sources

Can always decompose into isotropic + deviatoric

$$\underline{\underline{M}} = \frac{1}{3} (\text{tr } \underline{\underline{M}}) \underline{\underline{I}} + \underline{\underline{M}}^{\text{dev}}$$

ideal fault source in isotropic  $\oplus$ :  $\text{tr } \underline{\underline{M}} = 0$

$$\underline{\underline{M}} = \int_{\Sigma} \mu \Delta s_f (\hat{\underline{v}} \hat{\underline{v}} + \hat{\underline{v}} \hat{\underline{v}}) dA$$

curved fault

Common to impose linear constraint  $\text{tr } \underline{\underline{M}} = 0$   
no volume change  $\oplus$  source

Some evidence for  $\text{tr } \underline{M} = 0$  mechanisms in volcanic settings

CMT unknowns  $\rightarrow 9$

May also wish to compare with best double couple

det  $\underline{M}$  for a double couple

Non-linear: hard to impose

Instead CMT finds best-fitting double couple

$$(\underline{M}_{\text{bfde}} - \underline{M}) : (\underline{M}_{\text{bfde}} - \underline{M}) = \min$$

Solution found by diagonalization

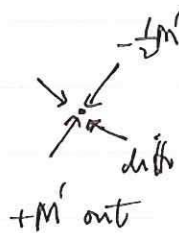
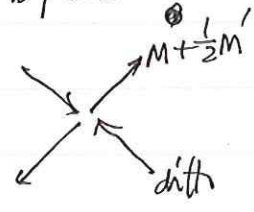
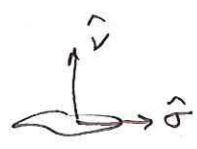
$$\underline{M} = \begin{pmatrix} M_{\text{bfde}} & & \\ & -M_{\text{bfde}} - M' & \\ & & M' \end{pmatrix} \quad \text{where } |M_{\text{bfde}}| \geq |M'|$$

note  $\text{tr } \underline{M} = 0$

$$= \begin{pmatrix} M_{\text{bfde}} + \frac{1}{2}M' & & \\ & -M - \frac{1}{2}M' & \\ & & 0 \end{pmatrix} + \begin{pmatrix} -\frac{1}{2}M' & & \\ & -\frac{1}{2}M' & \\ & & M' \end{pmatrix}$$

bfde

chd





CMT also gives so-called major and minor double couple

$$= \underbrace{\begin{pmatrix} M & & \\ & -M & \\ & & 0 \end{pmatrix}}_{\text{major}} + \underbrace{\begin{pmatrix} 0 & & \\ & -M' & \\ & & M' \end{pmatrix}}_{\text{minor}}$$

4% of events in CMT are significantly non-double-couple

~~Reasons~~ Reasons: ① noise in estimate  
② anisotropy

$$M = \int_{\Sigma} \underline{c} : \hat{\nu} \underline{AS}_f dA$$

probably small

③ curvature of fault surface

$$\underline{M} = \int_{\Sigma} \mu AS_f (\hat{\nu} \hat{\sigma} + \hat{\sigma} \hat{\nu})$$

Candidate geometries limited:

e.g.  $\hat{\sigma} = \text{const}$

need not be whole can



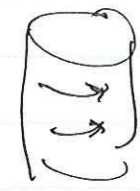
$$\underline{M} = \int_{\Sigma} \mu AS_f \hat{\nu} dA \hat{\sigma} + \hat{\sigma} \int_{\Sigma} \mu AS_f \hat{\nu} dA$$

$\hat{v}$  constant



can do same

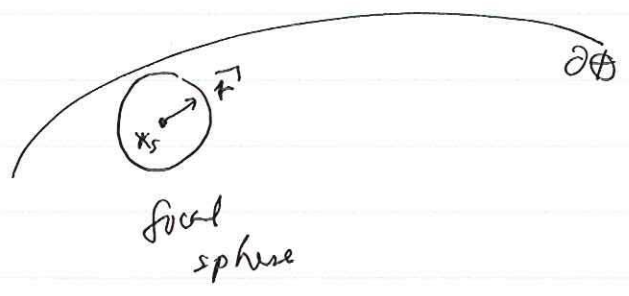
$\hat{v} \times \hat{z}$  constant



need not be whole can

Most interesting — example — Allan Ekström's volcanic ring faults.

Beach balls  $\hat{r} \cdot \underline{M} \cdot \hat{r}$



end class #9

NON-DOUBLE COUPLE COMPONENT

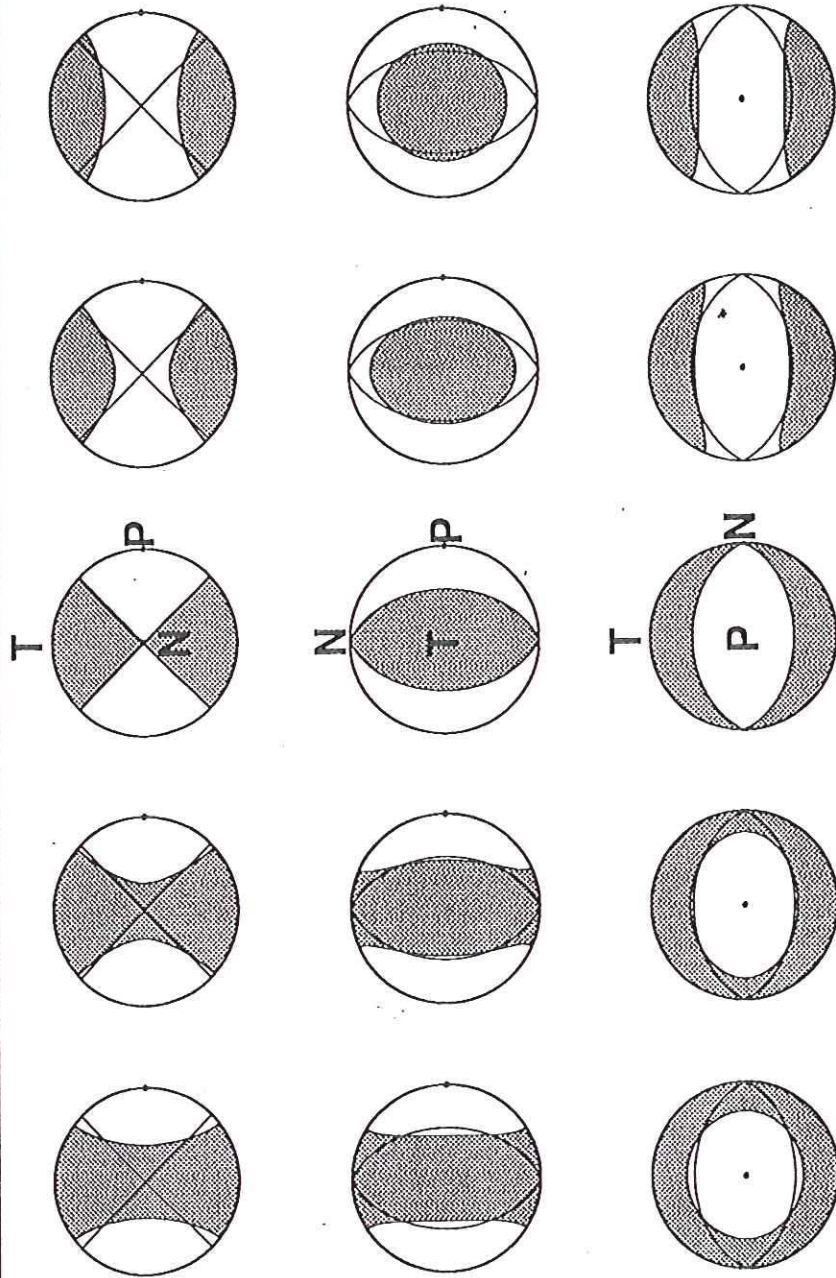
$\epsilon = -0.40$     $\epsilon = -0.20$     $\epsilon = 0.00$     $\epsilon = 0.20$     $\epsilon = 0.40$

*fault type*

*strike-slip*

*reverse*

*normal*

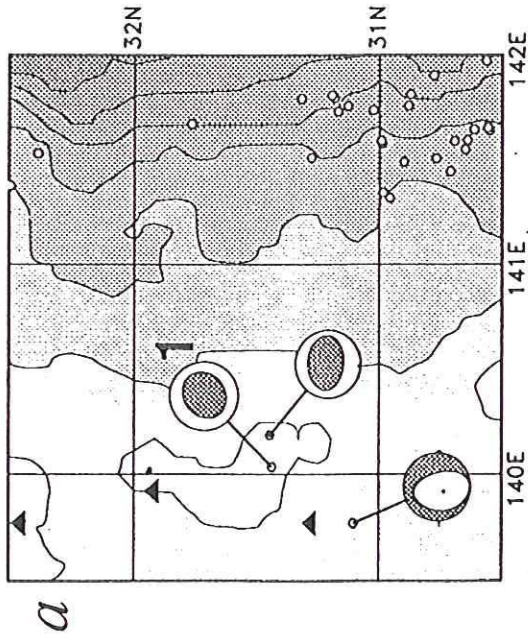


$$\epsilon = \frac{-M_2}{\max(M_1, -M_3)}$$

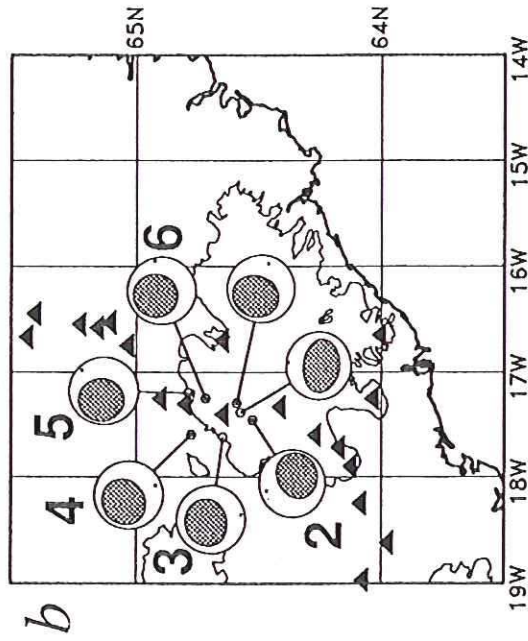
where  $M_1 > M_2 > M_3$



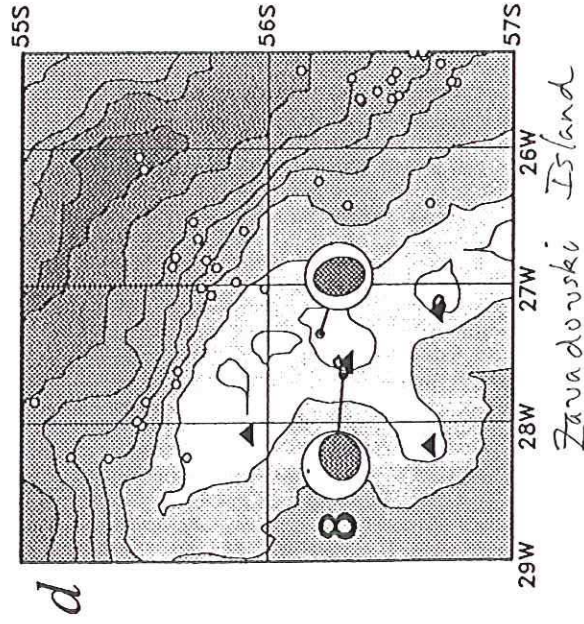
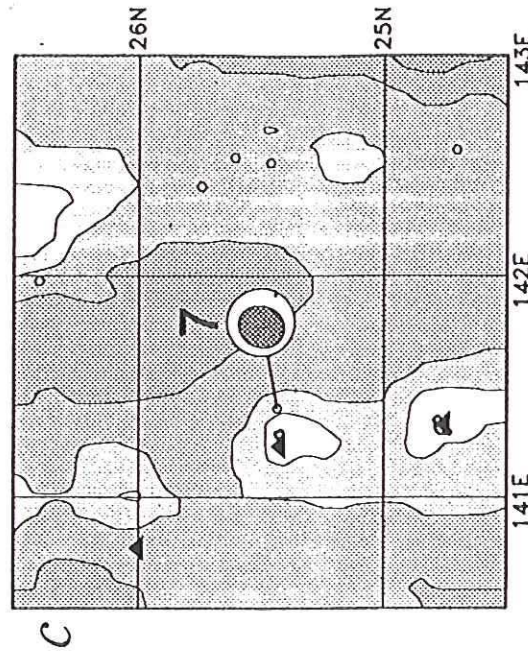
Tori Shima, Honshu



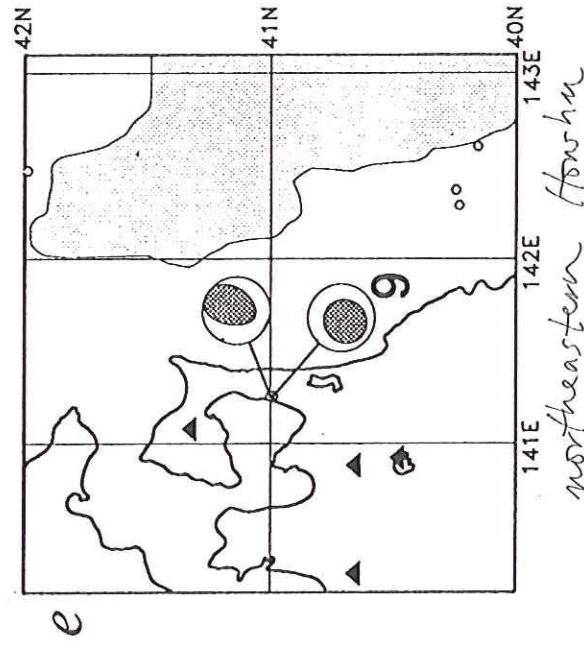
Bartharbunga volcano, Iceland



Volcan

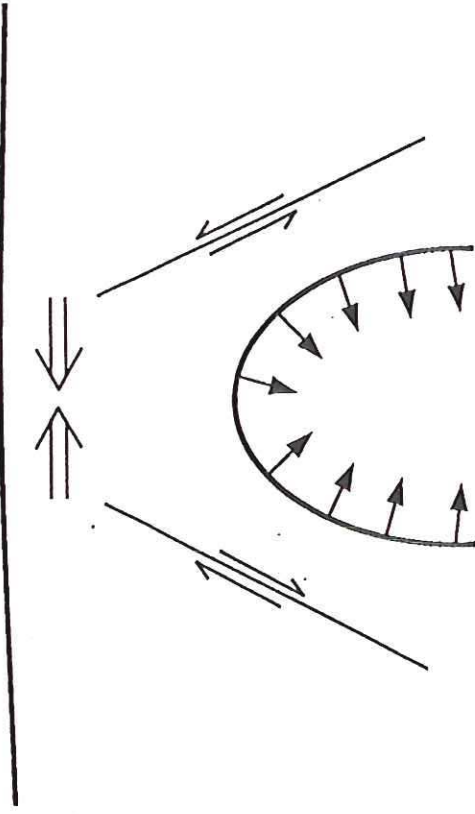


Zavardovski Island

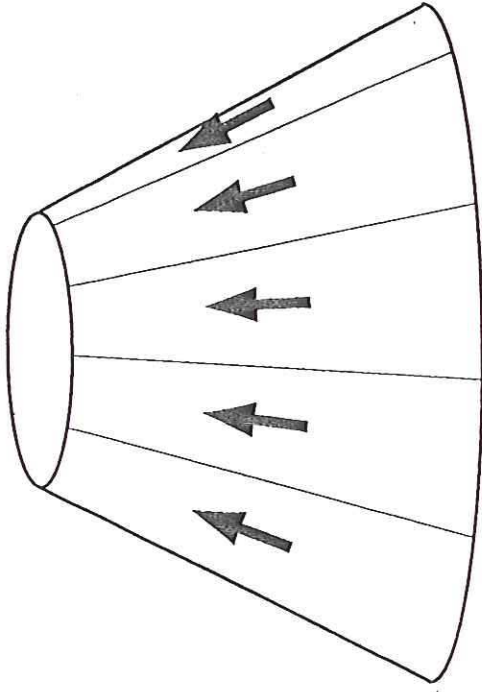


northeastern Honshu

a



b



volcanic ring fault



arc

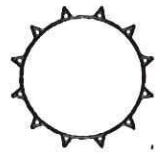
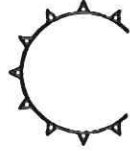
0°

90°

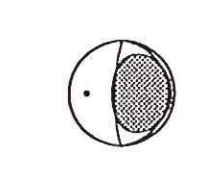
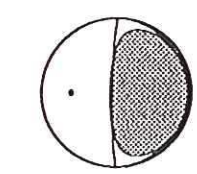
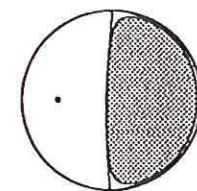
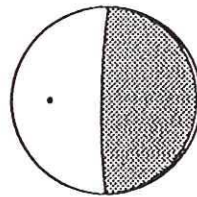
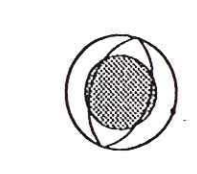
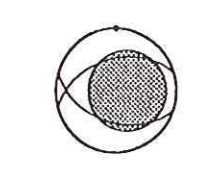
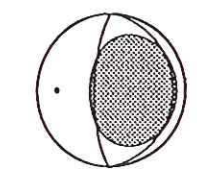
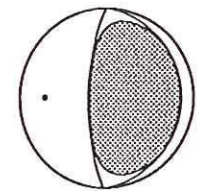
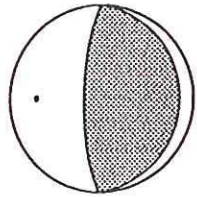
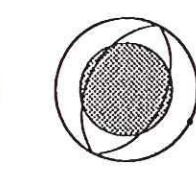
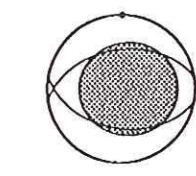
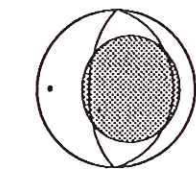
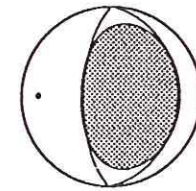
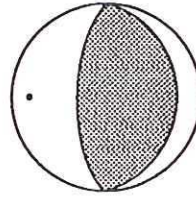
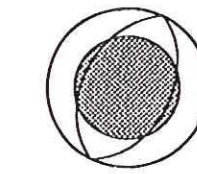
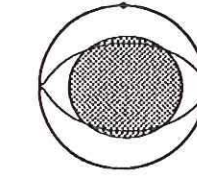
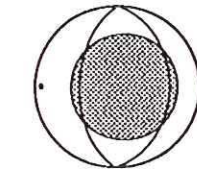
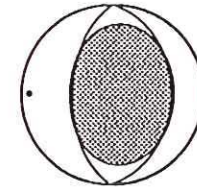
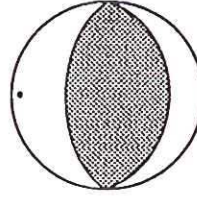
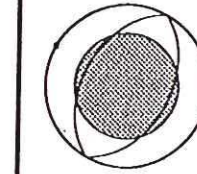
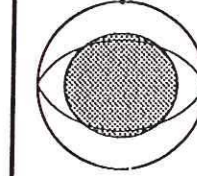
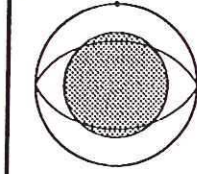
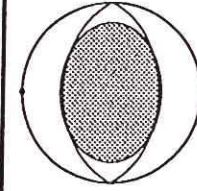
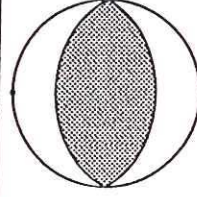
180°

270°

360°



dip



scalar moment  
 $\epsilon$